

DIGITAL MAPPING TECHNIQUES 2014

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University of Delaware

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> See Presentations and Proceedings from the DMT Meetings (1997-2014) http://ngmdb.usgs.gov/info/dmt/



Impacts of sea-level rise on groundwater resources in the Delaware coastal plain: a numerical model perspective

1. Abstract

Groundwater flow in Delaware's surficial aquifer adjacent to the Delaware Estuary is simulated using a 3-D, transient, variable density groundwater flow model. The model predicts movement of the fresh-water/salt-water interface and changes in water table depth due to sea-level rise through the year 2100. Three scenarios simulate sea level rises of 1.5 (S1), 1.0 (S2), and 0.5 (S3) meters. A representative conceptual model was constructed based on the characteristics of ten selected Delaware watersheds. Results indicate that the salt water intrudes inland up to 4.6 km along the tidal river in scenarios S1 and S2. To estimate effects on Delaware watersheds, modeled changes in water table depth are applied to 18 watersheds by mapping model coordinates to each watershed. Areas potentially impacted by sea level rise are identified by evaluating two critical depths to water, 0 and 0.5m, representing groundwater inundation (waterlogging) and effective rooting depths of major local crops. Land area impacted ranges from 60 hectares for scenario S3 with critical depth 0m to 18,500 hectares for scenario S1 with critical depth 0.5m. For scenario S3, there is minimal impact for the 0m condition (60 ha), but significant impact for the 0.5m condition (4,400 ha). There is 5-9 times more area impacted by waterlogging from a rising water table than from surface water inundation for all scenarios except scenario S3 with the 50cm condition where it is 38 times more area. Over 60% of the impacted area in all scenarios is cropland.



DNREC Sea Level Rise Scenarios 1.5 -High (Scenario 1) Intermediate (Scenario 2) Low (Scenario 3 3 05

2025 2050 2075

The study focuses on the effect of sea level rise on the water-table aquifers in Delaware Bay watersheds

SEAWAT simulation code

3-D. transient, variable density

25 m/da

Hydrologic Properties

Model Parameter Cell size

ongitudinal dispe fective porosity

Initial condition:

1.000-vr ramp-up to steady-state

Hydraulic Col

Numerical Model Grid

2. Modeling Methodology Conceptual model of a typical tidal creek and Conceptual Model (simple) watershed in the Delaware Estuary.

Conceptual model is a rectangular domain with an upland (gray), tidal river (blue), bay (turquoise) and, tidal wetlands along the river and bay (purple). As sea level rises, the creek stage increases with time to account for the effect of the encroaching tidal prism. The marsh area is progressively inundated during sea level rise to simulate transgression. Saltwater migrates from the bay up the creek as sea level rises Concentual Model (detailed) <13.7 km Ť₫. no flow boundary 2 km 5 km 18 km salt water brackish wat fresh water cross section(A-A

3.1. Water Table

At the unland/marsh boundary (X=2 000 m), the water table rises 0.8, 0.3 and 0.1 m for S1, S2, S3, respectively by year 2100. Head changes propagate 13, 8, and <1 kilometers inland for S1, S2, and S3, respectively by year 2100.

Contour maps of head change in the water table aquifer due to sea level rise (year 2100)







3.2. Salinity

S1

S2

S3

210

Calculated salinity concentration along cross section A-A' (distance from river Δy = 1,000 m)



Calculated position of the salt water front in bottom layer for different distances from river (Ay)



By year 2100, in S1 and S2, the toe of the salt water front migrates 4.6 km in the aguifer under the river. The toe does not migrate significantly in S3 at that location. At 3.95 km from the river, the toe migrates 1.5, 1.5 and 0.5 km for S1, S2, and S3, respectively



4. Application

Information from the simulation was used to identify areas within these coastal watersheds that would be inundated by the rising sea or become waterlogged due to a rising water table.

4.1. Application Methodology

The predicted change in head (Δh) for each scenario was output from the model into a coordinate system representing distance from the present upland/marsh boundary (x), and distance from the river (y). A corresponding curvilinear coordinate system was d for each watershed. The GIS processing steps are given below



4.2. Application Results

A critical depth to water is defined as the depth where there would be impacts to current land uses. Two critical depths to water were used: 0 m (vater at land surface) and 0.5 m. The latter value was chosen as a conservative representation of the effective rooting depths (critical) folds of conservative and the surface) and 0.5 m. The latter value was chosen as a conservative representation of the effective rooting depths (critical) folds of conservative and the surface). The CPN of conservative representation of the effective rooting depths (critical) solutions of the depth or the conservative representation of the effective rooting depths (critical) solutions of the depth or the conservative representation of the effective rooting depth (critical) water depth or the solutions of conservative representations shown below, but total distributions of areas meeting the depth or their are shown below in blue (inundiated by surface value) and red (valetoging from rising water table). Most of the areas affected are corpland (c) eXPM of all scenarios and conditions, see (lect and photo below).







5. Conclusions

Water-table rise and surface-water inundation in year 2100

- Total land area impacted ranges from 60 hectares (ha) for 0.5m SLR with critical depth of 0m to 18,500 ha for 1.5m SLR with critical depth of 0.5m
- Over 60% of the area impacted in all scenarios is cropland.
- 3 to 9 times more area is impacted by a rising water table than from surface-water inundation for all scenarios except 0.5 m SLR with the 0.5m condition where it is 38 times more area

Salt-water intrusion in year 2100

• By year 2100, for 1.5m and 1.0m SLR, the salt water in the base of the aquifer migrates 4.6 km inland from the marsh/upland boundary under the

0.5m_gw 0m_gw inund j

· By year 2100, at 4 km from the river, salt water in the base of the aquifer migrates 1.5, 1.5 and 0.5 km for 1.5, 1.0, and 0.5m SLR, respectively.