U.S. DEPARTMENT OF THE INTERIOR
U.S. GEOLOGICAL SURVEY

SCIENTIFIC INVESTIGATIONS MAP 2812
Version 1.0

GEOLOGIC MAP OF THE WEST HALF OF THE KLAMATH FALLS 1° × 2° QUADRANGLE, SOUTH-CENTRAL OREGON

By
David R. Sherrod and Leda Beth G. Pickthorn

1992

CORRELATION OF MAP UNITS

Holocene
Pleistocene
Pliocene
Miocene

QUATERNARY
TERTIARY

Qa, Qaf, Qpf, Qs, Qg, Qf, Qls, Qv, Qba, Qa, Qd, Qas, Qr, QTs, Tbk, Tb1, Tb2, Tb3, Tv, Typt, Ti, Tcs, Tst, Tba, Ta, Tad, Td, Tr, Tt, Twt, Topt

GEOLOGIC SUMMARY

The west half of the Klamath Falls 1° by 2° quadrangle encompasses about 9,000 km², most of which is in the Basin and Range province of south-central Oregon (fig. 1). The map also includes a small part of the Cascade Range province along the eastern flank of the range. Bedrock in the map area is mostly upper Miocene and Pliocene basalt and basaltic andesite, but the rocks vary in composition from basalt to rhyolite. Continental fluviolacustrine strata locally interfinger with the volcanic rocks. In the north half of the map area, the entire sequence becomes younger to the west, culminating with Pleistocene and Holocene rocks and deposits in the Cascade Range. In the south half, structural complications and inadequate isotopic dating preclude specifying the overall age pattern of the sequence.

Basalt and minor basaltic andesite lava flows—the basalt of Basin and Range (Tb_1, Tb_2, and Tb_3)—predominate in all but the northwestern part of the map area. The flows were erupted from vents associated with (1) scattered shield volcanoes ranging from 2 to 4 km in diameter and as much as 500 m high, (2) tuff cones that formed where the vents erupted into lakes or through a zone of abundant ground water, and (3) numerous cinder cones and fissure vents. Many of the vents are along faults. These lava flows were erupted during the past 9 m.y. (Wells and Peck, 1961; Peterson, 1982; Hering, 1981; Pickthorn and Sherrod, in press). Please check with your BTR. range from 6.88 to 1.9 Ma, indicating the flows are mostly late Miocene and Pliocene (table 1).

Basaltic andesite (Tba) dominates at Yamsay Mountain, a broad volcano about 25 km long, 20 km wide, and about 1,000 m high, whose summit area includes a buried caldera and an arcuate segment of a ring fault (Hering, 1981). Fuego Mountain, 30 km to the south-southwest, is also built chiefly of basaltic andesite. Other areas mapped as basaltic andesite are scattered throughout the map area. The basaltic andesite, which is undated by radiometric methods, is late Miocene and Pliocene in age on the basis of its interfingering relations with adjacent dated volcanic rocks. Pleistocene basaltic andesite (Qba) is exposed adjacent to Mount Mazama (fig. 1, Crater Lake caldera) in the Cascade Range.

Pliocene andesite and dacite (Ta, Tad, and Td) include lava flows and domes on Yamsay Mountain (Hering, 1981), as well as previously unmapped andesite and dacite at Riverbed Butte, Bly Mountain, on the east flank of Chiloquin Ridge, in the center of Eagle Ridge (west of Upper Klamath Lake), and near Fort Klamath (fig. 1). All are slightly to moderately porphyritic and resemble calc-alkaline rocks of the Cascade Range; chemical analyses are available only for the Yamsay Mountain lava (Hering, 1981). Quaternary andesite and dacite (Qas, Qa, and Qd) are exposed in the Cascade Range on the east flank of Mount Mazama (northwest corner of map area).

The oldest silicic rocks are rhyodacite to rhyolite domes (Tr) and small-volume pyroclastic flows (Twt and Tt) on the east and west sides of Yamsay Mountain, at the Black Hills and Ferguson and Medicine Mountains, and near Bly Mountain. Extensive sheets of ash-flow tuff are lacking in this part of the Basin and Range province. Part of the Hager Mountain dome (fig. 1), which erupted mainly north of the map area (MacLeod and others, 1976), is exposed in the northeast corner of the map area. The domes have K-Ar ages ranging from 6.06±0.09 to 5.15±0.21 Ma (table 1), whereas the pyroclastic flows at Yamsay Mountain (Silver Creek Welded Tuff of Hering, 1981 and 1982) have K-Ar ages of 7.36±1.58 and 6.95±1.13 Ma. The only other silicic domes within the map area are in the Cascade Range near Mount Mazama and are middle Pleistocene in age; their radiometric ages range from 0.7 to 0.6 Ma (Bacon, 1983, addendum to p. 63). A 5.5-Ma age (table 1, map No. 13) obtained earlier by E.H. McKee (in Bacon, 1983) is now considered incorrect.

Continental sedimentary strata (Tcs)—diatomite, mudstone, sandstone, and conglomerate—were deposited during late Miocene and Pliocene time in a series of freshwater lakes and deltas in the central and southern parts of the map area (Newcomb, 1958). Pliocene diatomite also crops out west of the map area in the Cascade Range (Smith and others, 1982). Many of these rocks were previously assigned to the Yonna Formation by Newcomb (1958), the type section of which contains sedimentary (fluviolacustrine) strata overlain by basaltic lapilli tuff (palagonitic tuff cones). As originally mapped, the Yonna Formation also included large areas of primary volcaniclastic rocks such as silicic tuff (welded and unwelded pyroclastic flows) and lahar deposits. Because our mapping distinguishes volcanogenic deposits from continental sedimentary rocks, we here abandon the Yonna Formation, reassigning its rocks to several unnamed units (that is, Typt, Tcs, Tt, and Topt).

The unit of continental sedimentary rocks (Tcs) is latest late Miocene and Pliocene in age, on the basis of K-Ar ages (Pickthorn and Sherrod, in press). Near Beatty the continental sedimentary rocks are underlain by basalt lava flows, silicic domes, and pyroclastic flows that range in age from 6.88 to 5.52 Ma; the unit is overlain by 3.61-Ma basalt (table 1). Near Klamath Falls, a lava flow that has an age of 4.47±0.28 Ma is interbedded with the continental sedimentary rocks (table 1, map No. 9). Near Wildhorse Ridge (fig. 1), conglomerate and sandstone (Tcs) are overlain by 3.68-Ma basalt lava (table 1, map No. 8). Thus, most exposures of the unit of continental sedimentary rocks are probably between about 5.5 Ma and 3.6 Ma in age. This age assignment also applies to the lower part of the Yonna Formation, which on the basis of mammalian and fish fossils was considered Pliocene in age (broadly 12 to 2 Ma for the time scale then in use) (Newcomb, 1958).

Sedimentation in the map area was greatly restricted during Pliestocene time, probably because the Klamath River had established its present-day course through the Cascade Range by late Pliocene time and integrated previously separate basins. Pleistocene sedimentary deposits include alluvium along valley floors, a few alluvial fans and talus cones, and graben-filling fluviolacustrine beds in the basins of Upper and Lower Klamath Lakes and the closed depression of Swan Lake Valley. Even today, lacustrine deposits, including diatomaceous strata, continue to form in Upper Klamath Lake, one of the largest freshwater lakes in the western United States. Lower Klamath Lake is now mostly agricultural land, reclaimed since 1940 by the U.S. Bureau of Reclamation.

STRUCTURE

North- to northwest-striking normal faults are the predominant structural feature in the west half of the Klamath Falls 1° by 2° quadrangle. Displacement is mostly less than 100 m in the central and eastern parts of the map area but increases to as much as 2 km on faults in the Klamath basin near the southwest corner of the map area. Exposed fault surfaces dip 50° to 60°, and most slickensides and mullions plunge directly downdip, consistent with normal slip.

The Klamath basin is a composite graben that forms the westernmost structural trough of the Basin and Range physiographic province. Extensional strain across the basin is probably less than 15 percent, although the estimates are necessarily conjectural because of uncertain fault geometry. Graben-bounding faults at the present-day level of erosion are broadly planar. East of the basin, decreasing number of faults and diminished displacement along them indicate greatly diminished strain (fig. 2).

In the southern part of the basin, a simple graben largely coincident with the original extent of Lower Klamath Lake coincides with a 20-mGal gravity depression (fig. 2; Sammel and Peterson, 1976; Couch and others, 1981). The gravity anomaly suggests that sediment as thick as 2 km (Sammel and Peterson, 1976) fills the graben. At Klamath Falls the graben shoals, and Quaternary sediment laps against upper Miocene and lower Pliocene lava flows and sedimentary rocks. In the northern part of the basin, Upper Klamath Lake floods, in part, a graben that reaches nearly to Crater Lake. The eastern escarpment of this graben has 450 m of structural relief in the Klamath Falls area. Although topographic relief diminishes northward, structural relief probably persists. Gravity data indicate that 0.4 to 1.2 km of sedimentary deposits (density contrast ranging from 1.0 to 0.3 gm/cm³, respectively) probably fill the basin north of Modoc Point (Veen, 1982).

The inception of Klamath graben faulting is poorly constrained. The Sprague River, which drains into the graben's north end, was probably established along its present-day route by 3 Ma. At about the same time, the basalt of Knot Tableland flowed south and west into the Sprague River valley. Presumably, downcutting at the mouth of the Sprague was controlled by both the nascent Klamath graben and the headward erosion of the Klamath River through the Cascade Range and into the upper Klamath basin.

Most of the Klamath graben is Pleistocene in age, however, and some faults have been active in Holocene time. The 400 m of topographic structural relief near Klamath Falls developed after eruption of lava flows as young as 2.80 to 1.9 Ma, which is the age of rocks capping the escarpment near Modoc Point (table 1, map N 1 and 3). Escarpments throughout the basin are commonly straight, steep, and relatively uneroded. Talus deposits exposed in the hanging wall of several youthful escarpments are themselves faulted and not merely scree buttressing against an exposed fault plane, as seen in excellent exposures where these deposits are quarried (for example, in Klamath Hills east of Lower Klamath Lake; near U.S. Highway 97 west of Lower Klamath Lake; and along U.S. Highway 97 east of Upper Klamath Lake). Upper Klamath Lake, which averages between 3 and 6 m deep, lacks alluvial fans along its escarpment-bounded margins, a condition most easily explained by continued Holocene subsidence. An elongate bathymetric trough ranging from 6 to 12 m deep and coincident with a fault on the west side of the lake is probably at least partly younger than basin-filling ash flows erupted from Mount Mazama about 6,845 ^{14}C yr B.P. (age from Bacon, 1983), on the assumption that pyroclastic debris would have choked and buried the narrow trough.

Northeast-striking faults are rare. Many closely spaced, short northeast-trending lineaments of uncertain origin are shown west of Yamsay Mountain (north-central part of map), in an area blanketed by Holocene pumice-fall deposits (Qpf) and underlain by relatively easily eroded sedimentary and pyroclastic rocks (Tcs and Ts).

FOLDS

Two folds shown in Yonna Valley are inferred from the outcrop pattern of map units and dip of bedding. They may result from extensional processes such as reverse and normal drag flexures (for example, Hamblin, 1965) or perhaps from localized contraction due to differential subsidence. The interpretation of folding could be incorrect, and the oppositely dipping beds may indicate differentially tilted fault blocks. Exposures in Yonna Valley are too poor to resolve this uncertainty.

Folded near-vent deposits are exposed in a roadcut along U.S. Highway 97 at Rattlesnake Point on the east side of Upper Klamath Lake (fig. 1), where a northwest-striking fault displaces the deposits against an adjacent block to the west. The fold, which cannot be traced beyond the outcrop and is not shown on the map, is an asymmetric anticline that may have resulted from drag created during oblique slip. No slickensides are exposed on the fault plane.

REFERENCES CITED

Bacon, C.R., 1983, Eruptive history of Mount Mazama and Crater Lake caldera, Cascade Range, U.S.A.: Journal of Volcanology and Geothermal Research, v. 18, p. 57–115.

Couch, R.W., Pitts, G.S., Veen, C.A., and Gemperle, Michael, 1981, Free-air gravity anomaly map and complete Bouguer gravity anomaly map, Cascade Mountain Range, southern Oregon: Oregon Department of Geology and Mineral Industries Geological Map Series GMS–16, scale 1:250,000.

Flebelkorn, R.B., Walker, G.W., MacLeod, N.S., McKee, E.H., and Smith, J.G., 1983, Index to K-Ar age determinations for the state of Oregon: Isochron/West, no. 37, p. 2–60.

Hamblin, W.K., 1965, Origin of reverse drag on the downthrown side of normal faults: Geological Society of America Bulletin, v. 76, no. 10, p. 1145–1164.

Hering, C.W., 1981, Geology and petrology of the Yamsay Mountain complex, south-central Oregon: A study of bimodal volcanism: Eugene, University of Oregon, Ph.D. dissertation, 194 p.

——1982, Geology and petrology of the Yamsay Mountain complex, south-central Oregon: A study of bimodal volcanism: Dissertation Abstracts International, v. 42, no. 11, p. 4332–B.

Johnson, D.M., Petersen, R.R., Lycan, D.R., Sweet, J.W., Neuhaus, M.E., and Schaedel, A.L., 1985, Atlas of Oregon Lakes: Corvallis, Oregon State University Press, 317 p.

MacLeod, N.S., Walker, G.W., and McKee, E.H., 1976, Geothermal significance of eastward increase in age of upper Cenozoic rhyolitic domes in southeastern Oregon, *in* Second United Nations Symposium on the Development and Use of Geothermal Resource, San Francisco, Calif., May 20–29, 1975, Proceedings, v. 1: Washington D.C., U.S. Government Printing Office, p. 465–474.

McKee, E.H., MacLeod, N.S., and Walker, G.W., 1976, Potassium-argon ages of Late Cenozoic silicic volcanic rocks, southeastern Oregon: Isochron/West, no. 15, p. 37–41.

McKee, E.H., Duffield, W.A., and Stern, R.J., 1983, Late Miocene and early Pliocene basaltic rocks and their implications for crustal structure, northeastern California and south-central Oregon: Geological Society of America Bulletin, v. 94, no. 2, p. 292–304.

Mereweather, E.A., 1953, The geology of the lower Sprague River area, Klamath County, Oregon: Eugene, University of Oregon, M.S. thesis, 62 p.

Newcomb, R.C., 1958, Yonna Formation of the Klamath River Basin, Oregon: Northwest Science, v. 32, no. 2, p. 41–48.

Peterson, N.V., and McIntyre, J.R., 1970, The reconnaissance geology and mineral resources of eastern Klamath County and western Lake County, Oregon: Oregon Department of Geology and Mineral Industries Bulletin 66, 70 p.

Pickthorn, L.B.G., and Sherrod, D.R., in press, Potassium-argon ages from Klamath Falls area, south-central Oregon: Isochron/West.

Sammel, E.A., and Peterson, D.L., 1976, Hydrologic reconnaissance of the geothermal area near Klamath Falls, Oregon, with a preliminary interpretation of geophysical data: U.S. Geological Survey Water Resources Investigation Open-File Report WRI 76-127, 129 p.

Smith, J.G., Page, N.J, Johnson, M.G., Moring, B.C., and Gray, Floyd, 1982, Preliminary geologic map of the Medford 1° by 2° quadrangle, Oregon and California: U.S. Geological Survey Open-File Report 82–955, scale 1:250,000.

Steiger, R.H., and Jäger, E., 1977, Subcommission on geochronology: Convention on the use of decay constants in geo- and cosmochronology: Earth and Planetary Science Letters, v. 36, no. 3, p. 359–362.

Veen, C.A., 1982, Gravity anomalies and their structural implications for the southern Oregon Cascade Mountains and adjoining Basin and Range province: Eugene, University of Oregon, M.S. thesis, 86 p.

Wells, F.G., and Peck, D.L., 1961, Geologic map of Oregon west of the 121st meridian: U.S. Geological Survey Miscellaneous Geologic Investigations Map I–325, scale 1:500,000.

Williams, Howel, 1942, The geology of Crater Lake National Park, Oregon, with a reconnaissance of the Cascade Range southward to Mount Shasta: Carnegie Institute of Washington Publication 540, 162 p.

Table 1. *Summary of K-Ar ages in map area and vicinity*

Map No.	Sample No.	Location: Latitude (N.)	Location: Longitude (W.)	General geographic local and map unit	Rock type	Material dated	Calculated age (Ma)[1]	Reference
1	KFE-11	42°21.0′	121°49.0′	Southeast of Modoc Point; Tb_1	Basalt	Whole rock	1.9±0.5	M. O'Brien and S. Benson, in Flebelkorn and others (1983)[2].
2	KFW-6	42°14.0′ (location not exact)	121°50.0′ (location not exact)	South of Upper Klamath Lake; map unit uncertain	Andesite, tuff	Whole rock	2.3±0.2	Do.
3	SS-1	42°23.8′	121°48.6′	Southeast of Modoc Point; Tb_1	Basalt	Whole rock	2.80±0.06	Pickthorn and Sherrod (in press).
4	SS-6	42°30.9′	121°18.9′	Knot Tableland; Tbk	Basalt	Whole rock	3.0±0.9	Do.
5	SS-3	42°39.7′	121°34.8′	North of Calimus Butte; Tb_3	Basalt	Whole rock	3.44±0.13	Do.
6	SS-8	42°33.3′	121°06.1′	East of Black Hills; Tb_2	Basalt	Whole rock	3.61±0.14	Do.
7	SS-4	42°33.6′	121°39.3′	S'Ochoits Canyon; Tcs	Pumiceous lahar	Plagioclase	3.62±0.12	Do.
8	SS-2	42°54.8′	121°31.2′	Wildhorse Ridge; Tb_3	Basalt	Whole rock	3.68±0.08	Do.
9	SS-12	42°12.6′	121°47.2′	Klamath Falls; lava flow in Tcs	Basalt	Plagioclase	4.47±0.28	Do.
10	SS-5	42°33.6′	121°40.0′	S'Ochoits Canyon; Tb_2	Basalt	Plagioclase	4.56±0.57	Do.
11	M4-131	42°56.6′	121°19.5′	Yamsay Mountain; Td	Dacite	Whole rock	4.79±0.17	MacLeod and others (1976); McKee and others (1976), Flebelkorn and others (1983).
12	M4-130	42°54.9′	121°08.5′	Partin Butte; Tr	Rhyolite	Whole rock	5.15±0.21	Do.
13	MS-45	42°56.9′	121°54.8′	East of Crater Lake; Qr	Rhyodacite	Plagioclase	5.5±0.5	Bacon (1983), E.H. McKee, (written common., 1986).
14	M3-106	42°35.6′	121°13.4′	Black Hills; Tr	Rhyolite	Plagioclase	5.52±0.5	MacLeod and others (1976), McKee and others (1976), Flebelkorn and others (1983).
15	SS-7	42°33.4′	121°05.9′	East of Black Hills; Tb_2	Basalt	Plagioclase	5.78±0.12	Pickthorn and Sherrod (in press).
16	MS-2	42°21.1′	121°23.9′	Bly Mountain; Tr	Rhyolite	Plagioclase	5.8±4	Flebelkorn and others (1983).
17	M3-61	43°00.6′	121°01.2′	Hager Mountain; Tr	Rhyolite	Obsidian	6.06±0.09	MacLeod and others (1976), McKee and others (1976), Flebelkorn and others (1983).
18	SS-9	42°20.0′	121°09.1′	East of Yainax Butte; Tb_2	Basalt	Whole rock	6.16±0.13	Pickthorn and Sherrod (in press).
19	SS-10	42°16.0′	121°20.7′	Southwest of Yainax Butte; Tb_2	Basalt	Plagioclase	6.88±0.60	Do.
20	M4-74	43°01.6′	121°06.9′	West of Silver Creek; Twt	Silicic ash-flow tuff	Plagioclase	6.95±1.13	McKee and others (1976), Flebelkorn and others (1983).
21	M4-73	43°05.2′	121°11.2′	West of Silver Creek; Twt	Silicic ash-flow tuff	Plagioclase	7.36±1.58	Do.

[1] K-Ar ages were calculated from data on references using the constants for the radioactive decay and abundance of ^{40}K recommended by the International Union of Geological Sciences Subcommission on Geochronology (Steiger and Jäger, 1977). These constants are: $\lambda_\beta = 0.580 \times 10^{-10}$ yr^{-1}, $\lambda_\beta + \lambda_{EC} = 4.962 \times 10^{-10}$ yr^{-1}, and $^{40}K/K_{total} = 1.167 \times 10^{-4}$ mol/mol.

[2] Using a preprint, Flebelkorn and others (1983) published the data and referenced it to Lawrence Berkeley Laboratory Report No. 11240 (1981), which never was published.

[3] Age of 5.5 Ma is incorrect. Subsequent determinations from domes in the same area are 0.7 to 0.6 Ma (Bacon, 1983, addendum to p. 63).

[4] Sample locality is outside boundary of geologic map. See figure 1 for location.

Figure 1. Index map showing selected geographic features, locations of K–Ar samples (solid circle, number refers to table 1), and area of report (shaded).

Figure 2. Tectonic sketch map.

PURPOSES OF STYLE GUIDE

(Additions and deletions were made to the original I-2182, a color map, for example purposes only)

Guidelines shown on this sheet (from the Cartographic Technical Standards of the U.S. Geological Survey, 1978, and recent updates) are provided as an aid in preparing your map or creating a style sheet with your own software. The sheet shows components that an average map might contain, but it cannot show all possibilities and all the technical standards. If you need additional information, please consult a map editor.

These guidelines are based on standards that cartographers follow routinely, but they should be used with good judgment and common sense. Spacing, leading, type sizes, and line weights can be increased or decreased depending on the size of features shown, complexity of the map, and overall sheet or plate size.

Two groups of type styles are used on USGS maps: (1) the serif group (such as Souvenir, Times Roman, or New Century Schoolbook) and (2) the sans serif group (such as Univers, Helvetica, or Optima). Serifs are the little feet at the ends of strokes of letters; they create a consistent horizontal direction at the ends of strokes. Sans serif type does not have the little feet and does not show the same contrast of thick and thin strokes that is found in serif type.

Souvenir and Univers fonts were used for this map, and the font size and leading reflect that. However, you may use other type styles so long as serif and sans serif styles are used as indicated on this style guide and the font size and leading are adjusted accordingly. For instance, if Times Roman is used, we suggest adjusting the font size and leading from 9/11 to 10/12 for legibility and a conservative use of space.

EXPLANATION

	Type style and font size
S-9	Serif—Showing font size in points
SI	Serif italic
SB	Serif bold
SBI	Serif bold italic
Sans	Sans serif
SansB	Sans serif bold
SansI	Sans serif italic
3 Ld	Leading (space between lines of type)—In points
12/15	Type size/type size leading—In this case, 12-point/12-point type plus 3-point leading
0.006 (0.15)	Measurement for spacing or lineweight—In inches (millimeters)
□	Em space—Used to show indents, square is point-size of type being used. An em in 9-point type is 8 points wide by 8 points high, and an em in 9-point type is 9 points wide by 9 points high
1 M	Long dash (1 em)—Option, Shift, dash on Microsoft Word

BRANCH OF PRINTING SIZES FOR MACHINE-FOLDED MAPS

Sheet size (in inches)	Maximum image size (in inches)
35 x 45	32.5 x 43
36 x 54	33.5 x 52
42 x 58	40 x 56

The absolute maximum image area available is 41 x 56 inches

ISSN 0-607-88586-1
9 780607 885864

Any use of trade names is for descriptive purposes only and does not imply endorsement by the U.S. Government

This map was produced on request, directly from digital files, on an electronic plotter

For sale by U.S. Geological Survey Information Services, Box 25286, Federal Center, Denver, CO 80225. 1-888-ASK-USGS

GIS files and a PDF file for this map are available online at http://pubs.usgs.gov/map/2004/2812

STYLE GUIDE
(VERSION 1.4)

[Previous versions prepared in the Western Publications Group]