# Changes in Stratigraphic Nomenclature by the U.S. Geological Survey, 1968

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CONTRIBUTIONS TO STRATIGRAPHY

GEOLOGICAL SURVEY BULLETIN 1294-A



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 $\mathbf{III}$ 

#### CONTRIBUTIONS TO STRATIGRAPHY

# CHANGES IN STRATIGRAPHIC NOMENCLATURE BY THE U.S. GEOLOGICAL SURVEY, 1968

By George V. Cohee, Robert G. Bates, and Wilna B. Wright

#### LISTING OF NOMENCLATURAL CHANGES

In the following table, stratigraphic names adopted, revised, reinstated, or abandoned are listed alphabetically. The age of the unit, the revision, and the area involved, along with the author's name and date of publication of the report, are given. The publications in which the changes in nomenclature were made are listed in the references at the end of this publication. The capitalization of age terms in the age column follows official usage.

One of the significant nomenclature changes in 1968 was the adoption by the U.S. Geological Survey of Holocene to replace Recent (Cohee, 1968). Holocene, meaning "wholly recent" and referring to the percentage of living organisms, originally was proposed as a "stage" following the Pleistocene "stage" by the Portuguese committee to the Third International Congress of 1885 (Morrison and others, 1957).

At the annual meeting of the American Commission on Stratigraphic Nomenclature, November 22, 1967, the Commission endorsed the use of Holocene instead of Recent and expressed the hope that the term "Holocene" would be adopted officially by various geological organizations. The formal term "Recent" was ambiguous in referring to sedimentary deposits, fossils, and present-day shells involving "Recent" or recent time. Holocene was given series rank equal to that of the Pleistocene because both vertebrate and invertebrate faunas reflect marked changes between Pleistocene and Holocene Epochs and the archaeological record provides means for subdividing Holocene deposits.

#### Changes in stratigraphic nomenclature

Name	Age	Location	Revision and reference
Ajibik Quartzite	middle Precambrian	Michigan	Ajibik Quartzite placed in the Menominee Group of the Animikie Series. Age changed from Precambrian to middle Precambrian. (Gair and Thaden, 1968.)
Alcova Limestone Member (of Chugwater Formation).	Triassic	Wyoming	Alcova Limestone Member of Chugwater Formation raised in rank to Alcova Limestone of Chugwater Group. (Piniringos. 1968.)
Altonian Substage	Pleistocene	Illinois and Wisconsin	Altonian Substage of Frye and Willman (1960) adopted as basal substage of the Wisconsinan Stage. (Frye and others, 1968.)
mation).			Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemaugh is con- sidered of group rank and consists of the Casselman and Glenshaw Formations, and the Ames is considered the top member of the Glenshaw Formation. (Roen and others, 1968.)
Ammonia Tanks Member (of Timber Mountain Tuff.)	early Pliocene	Nevada	Age changed from Pliocene to early Pliocene. (Kistler, 1968.)
Amsden Formation	Late Mississippian and Early and Middle Pennsylvanian.	Wyoming	Age changed from Early and Middle Pennsylvanian to Late Mississippian and Early and Middle Pennsylvanian. (Malloy, 1967.)
			Apache Canyon Formation of Tyrrell (1957) adopted. Overlies Willow Canyon Formation; underlies Shellenberger Canyon Formation (Finall this report p. A32)
canics).			A psey Conglomerate Member adopted, overlies Hells Half Acre Tuff Member and underlies andesite of Table Mountain. (Krieger, 1968a.)
Aquia Formation	. Paleocene	Maryland, Delaware, and Virginia.	Age changed from Eocene to Paleocene. (Hazel, 1968.)
		Colorado, Montana,	Aravaing Member adopted. (Krieger, 1968a.) Age changed from early and middle Miocene to early Miocene. (Sato and Denson, 1967.)
Bachelor Mountain Rhyolite	Oligocene	Colorado	Age changed from middle or late Tertiary to Oligocene. (Steven and others, 1967.)
Banner Formation	Late Mississippian (Meramec)	. Nevada	Banner Limestone of Granger and others (1957) adopted as
	natian).		Banner Formation. (Coats, 1969.)  Bardstown Member adopted as middle member of Drakes Formation in north-central Kentucky. (Peterson, this report, A36.)
Bates Mountain Tuff	Oligocene or Miocene	Nevada	Ratas Mountain Tuff adopted (Stawart and McKag 1968)
Bathtub Formation Bell Springs Member (of Nugget Sandstone)	Early Cretaceous Triassic(?)	Wyoming	Bathtub Formation adopted (Drewes, 1968.) Bell Springs Member adopted as basal member of Nugget Sandstone. (Pipiringos, 1968.)
Belted Range Tuff	late Miocene	Nevada	. Age changed from Miocene and Pliocene(?) to late Miocene.
Bentley Formation	Pleistocene	Texas	(Noble and others, 1968.)  Bentley Formation extended into Texas. (Wilson, 1967).

Bigford Member (of Mount Selman Formation) -	middle Eocene	Southern Texas	In the Rio Grande embayment south of Frio River, the Bigford Member of the Mount Selman Formation is raised to fornation rank, and the Mount Selman Forma- tion is abandoned. (Eargle, 1968.)
Big Rock Conglomerate Member (of Kiawa	Precambrian	New Mexico	Big Rock Conglomerate Member of Barker (1958) adopted.
			(Barker, this report, p. A21.) Birch Creek Schist restricted. Some of the rocks formerly assigned to the Birch Creek Schist now assigned to the Keevy Peak Formation. (Wahrhaftig, 1968.)
Birmingham Shale Member (of Conemaugh Formation).	Late Pennsylvanian	Pennsylvania and Ohio	Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemaugh is conconsidered of group rank and consists of the Casselman and Glenshaw Formations, and the Birmingham is con-
Bisbee Group	Early Cretaceous	Southeastern Arizona	sidered the basal member of the Casselman Formation. (Roen and others, 1968.) In the Empire and Santa Rita Mountains, the Bisbee Group includes (ascending order): Glance Conglomerate, Willow Canyon, Apache Canyon, Shellenberger Canyon, and Turney Ranch Formations (Finnell, this report, p. A28.)
Black Mingo Formation			Age changed from Eocene to Paleocene and Eocene. (Hazel,
Blanco Basin Formation	early Eocene	Southwestern Colorado	Age changed from Oligocene(?) to early Eocene. (Steven and others, 1967.)
			Indian Mills Sandstone Member of Reger and Price (1926) adopted as member of Bluefield Formation. (Englund,
Blue Mesa Tuff Bluestone Formation (of Pennington Group)	late OligoceneLate Mississippian and Early Pennsylvanian.	ColoradoVirginia and West Virginia	1968a.) Blue Mesa Tuff adopted. (Olson and others, 1968.) Formation divided into (ascending order): Pride Shale, Glady Fork Sandstone, gray, red, Brannwell, and upper members. Age changed from Late Mississippian to Late Mississippian and Early Pennsylvanian. (Englund, 1968a.)
Bossier Formation (of Cotton Valley Group)		Arkansas.	Subsurface. The Bossier is expanded to include all the dark-gray shale above the lower member of the Smackover Formation and below the Schuler Formation. (Dickinson, 1982)
Bouse Formation Bramwell Member (of Bluestone Formation) Brayman Shale	Pliocene Late Mississippian Late Silurian	Southwestern ArizonaVirginia and West Virginia New York	Bouse Formation adopted. (Metzger, 1968). Bramwell Member adopted. (Englund, 1968a.) Age changed from Silurian to Late Silurian. (Paylides and
Brightseat Formation	Paleocene	Maryland	Brightseat Formation made basal formation of Pamunkey
			Age changed from Cretaceous or early Tertiary to Late
			Cretaceous. (Hayes and Drewes, 1968.) Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemaugh is considered of group rank and consists of the Casselman and Glenshaw Formations, and the Brush Creek is
Brynt Draw Member (of Popo Agie Formation)	Late Triassic	Wyoming	considered a member of the Glenshaw Formation. (Roen and others, 1968.)  Brynt Draw adopted as basal member of Popo Agle Formation. (Pipiringos, 1968.)

Name	Age	Location	Revision and reference
Buffalo Sandstone Member (of Conemaugh Formation).	Late Pennsylvanian	Pennsylvania and Ohio	Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemaugh is con- sidered of group rank and consists of the Casselman and Glenshaw Formations, and the Buffalo is considered a member of the Glenshaw Formation. (Roen and others,
Bull Ridge Member (of Madison Limestone)	Forly Mississippion (souly	West southed West	1988 )
Produce Andrews	Meramec).	west-central wyoming	Bull Ridge adopted as upper member of the Madison Limestone. (Sando. 1968.)
Bulker Andesite	_ Oligocene	Southern Colorado	Limestone. (Sando, 1968.)  Name changed to Bunker Trachyandesite; age changed
Burnett Formation (of Puget Group)	. Eocene	Washington	from Eocene to Oligocene. (Steven and Epis, 1968.)  Name abandoned. Lower part now assigned to Carbonado  Formation and upper part to Spiketon Formation. (Gard, 1968.)
Burns Formation (of Silverton Volcanic Group)	Olimonana	Calanada	(Gald, 1908.)
Bursum Formation (of Magdalena Group)	Early Permian	New Mexico	Age changed from middle and late Tertiary to Oligocene. (Luedke and Burbank, 1968.) Formation restricted, geographically, to vicinity of Oscura Mountains, central New Mexico. (Bachman, 1968.) California Creek Member adopted. (Wahrhaftig, 1968.)
California Creek Member (of Totatlanika Schist).	Mississippian(?)	Alaska	California Creek Member adopted. (Wahrhaftig, 1968.)
Camarones Sandstone Cambridge Limestone Member (of Conemaugh Formation).	Late Cretaceous Late Pennsylvanian	Puerto Rico	Camarones Sandstone adopted. (Pease, 1968a.) Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemaugh is con- sidered of group rank and consists of the Casselman and Glenshaw Formations, and the Cambridge is considered a member of the Glenshaw Formation. (Roen and others, 1968.)
Camden Chert	Early Devonian	Kentucky, Tennessee, and	A mar of the control
Campbell Mountain Member (of Bachelor Mountain Rhyolite).	Oligocene	Colorado	Age changed from Early and Middle Devonian to Early Devonian. (Boucot and Johnson, 1968.)  Age changed from middle or late Tertiary to Oligocene. (Stayen, and others 1967.)
Cancel Breccia Canelo Hills Volcanics	Early to Late Cretaceous	Puerto Rico	(Steven and others, 1967.) Cancel Breccia adopted. (Pease, 1968a.) Age changed from Triassic and Jurassic to Late Triassic
Control II II	Jurassic.	Southeastern Arizona	Age changed from Triassic and Jurassic to Late Triassic and Early Jurassic. (Haves and Drewes, 1968.)
Cantwell Formation	Paleocene	Alaska	Age changed from Triassic and Jurassic to Late Triassic and Early Jurassic. (Hayes and Drewes, 1968.)  Age changed from Early Cretaceous to Paleocene. (Wolfe
Carbonado Formation (of Puget Group)	_ middle(?) Eocene	. Washington	and wannates, this report, p. A41.)  Carbonado Formation revised to include the Wilkeson  Formation (abandoned) and the lower part of the Burnett  Formation (abandoned). Age changed from Eocene to
Cardiff Conglomerate	Ordovician(?)	Maryland and Pennsylvania	Cardiff Conglomerate changed to Cardiff Metaconglom-
Carpenter Ridge Tuff	Oligocene	Colorado	erate. (Southwick and Fisher, 1967.) Carpenter Ridge Tuff adopted. (Olson and others, 1968.)
Cash Creek Quartzite.	Late Cretaceous	Puerto Rico	erate. (Southwick and Fisher, 1967.) Carpenter Ridge Tuff adopted. (Olson and others, 1968.) Carrafzo Breccia adopted. (Pease, 1968a.) Cash Creek Quartzite adopted. (Hobbs and others, 1968.)
• • • • • • • • • • • • • • • • • • • •	or made order form	144110	Cash Oreek Quartzite adopted. (110008 and others, 1968.)

Casper Formation	Early, Middle, and Late Pennsylvanian and Early Permian.	Southern Wyoming	Age changed from Pennsylvanian and Permian to Early Pennsylvanian through Early Permian. (Mallory, 1967.)
	. Late Pennsylvanian		Casselman Formation of Flint (1965) adopted wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized. In those areas, the Conemaugh is considered a group and includes the Glenshaw and Casselman Formations. Members of the Casselman (in ascending order) are: Birmingham Shale, Elk Lick Limestone, Morgantown Sandstone, Clarksburg Limestone, Connellsville Sandstone, Summerhill Sandstone, and Wilmore Sandstone. (Roen and others, 1968.)
			Cerro Gordo Formation of Lidiak (1965) adopted as Cerro Gordo Lava. (Pease, 1968a.)
			Chesterfield Range Group extended into northeastern Utah. (Sando, 1967.)
Chinquapin Metabasalt Member (of South Fork Mountain Schist).	Late(?) Cretaceous	Northwestern California	Chinquapin Metabasalt Member adopted. (Blake and
Chocolay Group (of Animikie Series)	middle Precambrian	Michigan	In the Marquette area, the Chocolay Group includes (in ascending order): Enchantment Lake Formation (new), Mesnard Quartzite, Kona Dolomite, and Wewe Slate. (Gair and Thaden, 1968.)
			Chugwater Formation raised to group rank; includes (in ascending order): Red Peak Formation, Alcova Limestone, Crow Mountain Sandstone (locally Jelm Formation), and Pone Agie Formation (Pinigingos, 1968)
Cipao Formation	. Tage Ongoverse and early	I UCITO INICO	Chute Creek Member adopted. (Wahrhaftig, 1968.)  Age changed from Oligocene and Miocene to late Oligocene
	middle Eocene		Cook Mountain Formation in Texas restricted to eastern and central parts; replaced by Laredo Formation in Rio Grande embayment south of Frio River. (Fargle, 1968.)
			Clamgulchian Stage (Floral) adopted. (Wolfe and others,
Formation).			Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemaugh is con- sidered of group rank and consists of the Casselman and Glenshaw Formations, and the Clarksburg is considered a member of the Casselman Formation. (Roen and others, 1968.)
Clayton Mine Quartzite Clear Creek Formation	Middle Ordovician or older Early Devonian	Idaho	Clayton Mine Quartzite adopted. (Hobbs and others, 1968.) Age changed from Early and Middle Devonian to Early Devonian. (Boucot and Johnson, 1968.)
Clyde Formation	Early Permian (Leonard)	Texas	Devonian. (Boucot and Johnson, 1968.)  Age changed from Early Permian (Leonard?) to Early Permian (Leonard). (Myers, 1968.)
			In the southwestern part of the North Park area, the Middle Park Formation is reduced in rank and made the basel member of the Coalmont Formation (Hall, 1968.)
Cockeysville Marble (of Glenarm Series)	late Precambrian(?)	Maryland, Pennsylvania, and	Age changed from early Paleozoic (?) to late Precambrian (?)
Compeau Creek Gneiss	early Precambrian	Michigan	(Southwick and Fisher, 1967.) Compeau Creek Gneiss adopted. (Gair and Thaden, 1968.)

Name	Age	Location	Revision and reference
Conejos Quartz Latite	Oligocene or older	Colorado	Age changed from middle or late Tertiary to Oligocene or
			older. (Olson and others, 1968.) Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemaugh is considered of group rank and includes two formations, the Glenshaw (basal) and Casselman Formations. The Ames is the top member of the Glenshaw. (Roen and others, 1968.)
Conestoga Limestone	Early Ordovician	Pennsylvania	Age changed from Late Cambrian and Early Ordovician to Early Ordovician. (Meisler and Becher, 1968.)
Formation)			Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemangh is considered of group rank and consists of the Casselman and Glenshaw Formations, and the Connellsville is considered a member of the Casselman Formation. (Roen and others,
Continental Granodiorite	Precambrian	Southeastern Arizona	Continental Granodiorite adopted. (Drewes, 1968.)  Cook Mountain Formation in Texas restricted to eastern
			and central parts; replaced by Laredo Formation in Rio
Copper Creek Granodiorite	Late Cretaceous and (or)	Arizona	Age changed from Tata(2) Chatagonia on Contions to Tata
			Cretaceous and (or) early Tertiary. (Krieger, 1968b.)  Lulbegrud Clay, Waco Limestone, and Estill Clay, all of Foerste (1905), adopted. Formation includes (in ascending order): Plum Creek, Oldham, Lulbegrud Shale, Waco, and Estill Shale Members. (Simmons, 1967.)
			Age changed from middle or late Tertiary to Oligocene.
Crow Mountain Sandstone Member (of Chugwater Formation).	Late Triassic	Wyoming	Crow Mountain Sandstone Member raised to formation
Darwin Sandstone Member (of Amsden Formation).	Late Mississippian	do	rank. Included in Chugwater Group. (Pipiringos, 1968.) Age changed from Early Pennsylvanian to Late Mississip-
Devils Hole Formation	late Eocene	South-central Colorado	pian. (Mallory, 1967.)  Age changed from Miocene(?) to late Eocene. (Steven and
Dewitt Formation	Miocene and Pliocene	Eastern Texas	Epis, 1968.)  Formation was named by Deussen (1914). In subsequent work by Deussen the name was not used nor was it ever
Diamond Peak Formation	late Mississippian and Early Pennsylvanian.	Nevada	used by any other author. Abandoned for nonuse.  Age changed from Mississippian to Late Mississippian and Early Pennsylvanian in the Carlin-Pinon Range area.
			(Smith and Ketner, 1968.) Dillon Mesa Tuff adopted. (Olson and others, 1968.) Dixville Formation of Green (1964) adopted. (Harwood and Berry, 1967.)
Drakes Formation	Late Ordovician (Cincinnatian).	North-central Kentucky	Bardstown Member adopted as middle member of Drakes Formation in north-central Kentucky. Overlies Rowland

Member and underlies Saluda Dolomite Member. (Peter-

			son, this report, p. A36.)
Eagle Valley Evaporite	Pennsylvanian and Permian	Colorado	Name changed from Eagle Valley Evaporite to Eagle Valley Formation. (Bartleson and others, 1968.)
Elephant Head Quartz Monzonite	Late Cretaceous.	Southeastern Arizona	Elephant Head Quartz Monzonite adopted. (Drewes, 1968.)
Elk Lick Limestone Member (of Conemaugh	Late Pennsylvanian	Pennsylvania and Ohio	Wherever the Ames Limestone Member (of the Glenshaw
Formation).			Formation) can be recognized, the Conemaugh is considered of group rank and consists of the Casselman and
			Glenshaw Formations, and the Elk Lick is considered a
			member of the Casselman Formation. (Roen and others,
Tille Delemite	Middle Ondtetem	T3.1.	1968.) Ella Dolomite adopted. (Hobbs and others, 1968.)
Ellensburg Formation	late Miocene and early	Oregon	Ellensburg Formation extended from Washington into
-	Pliocene	<del>.</del>	north-central Oregon (Weters 1968)
El Ocho Formation	Early to Late Cretaceous	Puerto Rico	El Ocho Formation adopted. (Pease, 1968a.) El Pico Clay adopted for the former upper unnamed mem-
El Pico Ciay (of Chamborne Group)	madie Eocene	Southern Texas	ber of the Mount Selman Formation (abandoned) in the
			Rio Grande embayment south of Frio River. (Eargle,
The best west Take Homestine (of Charles	and date Theorem bades	351-1-1	1968.)
			Enchantment Lake Formation adopted. (Gair and Thaden, 1968.)
Esch Creek Glaciation	Holocene	Alaska	1968.) Esch Creek Glaciation adopted. (Sainsbury, 1967.) Estill Clay of Foerste (1905) adopted as Estill Shale Mem-
Estill Shale Member (of Crab Orchard Formation).	Middle Silurian	Kentucky	Estill Clay of Foerste (1905) adopted as Estill Shale Member. (Simmons, 1967.)
Eureka Tuff (of Silverton Volcanic Group)	Oligocene	Colorado	Age changed from middle and late Tertiary to Oligocene.
			(Luedke and Burbank, 1968.)
Ewing Limestone Member (of Conemaugh Formation).	Late Pennsylvanian	Pennsylvania and Ohio	Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemaugh is con-
Formation).			sidered of group rank and consists of the Casselman and
			Glenshaw Formations, and the Ewing is considered a
			member of the Glenshaw Formation. (Roen and others, 1968.)
Fajardo Formation	Late Cretaceous	Puerto Rico	Fajardo Formation replaced by Rio Piedras Siltstone in
Table Mills Condition Manager (of Til to	Trade After the trade	***	Naranjito quadrangle. (Pease, 1968b.)
Formation).	Late Mississippian	Virginia and West Virginia	Falls Mills Sandstone Member of Reger and Price (1926) adopted. (Englund, 1968a.)
Farisita Conglomerate	late Eocene	South-central Colorado	Age changed from Oligocene(?) to late Eocene. (Steven and
Form dala Cilt	loto Disistacene	This sie on 3 Witnesselle	Epis, 1968.) Farmdale Silt of Frye and Willman (1960) adopted. (Frye
			and others, 1968.)
Farmdalian Substage (of Wisconsinan Stage)	do	do	Farmdalian Substage of Frye and William (1960) adopted.
			Overlies Altonian Substage and underlies Woodfordian Substage. (Frye and others, 1968.)
Farmers Creek Rhyolite	Oligocene	Colorado	Age changed from middle or late Tertiary to Oligocene.
			(Steven and Others, 1967.)
Figuera Voicanics	. iate Paleocene or early Eocene.	Puerto Rico	Figuera Volcanics abandoned; replaced by Guaracanal
Fish Canyon Tuff	late Oligocene	Colorado	Andesite. (Pease, 1968a.) Fish Canyon Tuff adopted. (Olson and others, 1968.)
Fisher Quartz Latite	Oligocene	do	Age changed from middle or late Tertiary to Oligocene.
Fleming Formation	Miocene	Louisians and Tayes	(Steven and others, 1967.) Fleming Formation, previously abandoned, reinstated.
		Tronsporter etter 1 page	(Anders and others, 1968.)

Name	Age	Location	Revision and reference
			Fort Crittenden Formation of Stoyanow (1949) adopted. (Drewes, 1968.)
Fraction Tuff	middle Miocene	Nevada	Age changed from late Miocene to middle Miocene. (Ander-
			son and Ekren, 1968.) The Frailes Formation, as used by Kaye (1959), is not a mappable unit in the report area and is included in the revised Guaynabo Formation, except for La Muda Limestone Member which is raised to formation rank. (Pease, 1968a.)
Franklinian Stage (Floral) Fultonian Stage (Floral) Galluro Volcanics	early Eocene middle Eocene Miocene.	West-central WashingtondoArizona.	Franklinian Stage (Floral) adopted. (Wolfe, 1968.) Fultonian Stage (Floral) adopted. (Wolfe, 1968.) Includes (in ascending order): Holy Joe, Aravaipa, Hells Half Acre Tuff, and Apsey Conglomerate Members. Age changed from middle Tertiary or younger to Miocene. (Krieger, 1968b.)
			Grove Creek Formation reduced to member rank and made upper member of Snowy Range Formation. Group includes (in ascending order): Pilgrim Limestone and Snowy Range Formation. (Pilgrey and Malon, 1982)
Formation) Grandstone Member (of Bluestone	Late Mississippian	Virginia and West Virginia	Gardner Canyon Formation adopted. (Drewes, 1968.) Ghost Rocks Formation adopted. (Moore, 1969.) Glady Fork Sandstone Member of Reger and Price (1926) adopted. (Englund, 1968a.)
Glenarm Series	late Precambrian(?)	Maryland, Pennsylvania, New Jersey, Delaware, and Virginia.	Age changed from early Paleozoic(?) to late Precambrian(?). In Maryland the following units are excluded from the Glenarm Series: Jamsville Phyllite, Marburg Schist, Silver Rum Limestone, Urbana Phyllite, and Wakefield Marble. In Maryland the following units are included in the Glenarm Series: Setters Formation, Cockeysville Marble, and the Wissahickon Formation (revised). (Southwick and Fisher, 1967.)
Glenshaw Formation (of Conemaugh Group)	Late Pennsylvanian	Pennsylvania and Ohio	Glenshaw Formation of Flint (1965) adopted wherever the Ames Limestone Member (of the Glenshaw Formation) ean be recognized. In those areas, the Conemagh is considered a group and includes the Glenshaw and Casselman Formations. Members of the Glenshaw (in ascending order) are: Uffington Shale, Mahoming Sandstone, Brush Creek Limestone, Buffalo Sandstone, Cambridge Limestone, Ewing Limestone, Saltsburg Sandstone, and Ames Limestone. (Roen and others, 1968.)
Glory Hole Volcanics	Late Cretaceous and (or)	Arizona	Age changed from Late(?) Cretaceous to early Tertiary to
Grape Creek Limestone Member (of Clyde Formation).	early Tertiary. Early Permian (Leonard)		Late Cretaceous and (or) early Tertiary. (Krieger, 1968b.) Age changed from Early Permian (Leonard?) to Early Permian (Leonard). (Myers, 1968.)

Gringo Gulch Volcanics	Paleocene(?)	Southeastern Arizona	Gringo Gulch Volcanics adopted. (Drewes, 1968.)
Grossman Formation	late(?) Oligocope	Southeastern Arizona	Grosvenor Hills Volcanics adopted. (Drewes, 1968.)
Groupe Convon Mamber (of Rolled Range Tuff)	late Miocene	Nevada	Age changed from Miocene and Pliocenc(?) to late Miocene
Ground Charlet and Control Poster Printing 2 and 2			(Noble and others, 1968.)
Grouse Canyon Member (of Indian Trail Formation).	do	do	(Noble and others, 1968.)  Age changed from Miocene and Pliocene(?) to Miocene. (Kistler, 1968.)
Grove Creek Formation	Late Cambrian	Wyoming and Montana	(Kistler, 1968.) Grove Creek Formation reduced to member rank and
			made upper member of Snowy Range Formation. (Pierce and Nelson, 1968.)
			Guaracanal Andesite adopted; replaces Figuera Volcanics which was abandoned. (Pease, 1968a.)
Guaynabo Formation	Late Cretaceous	do	Redefined to include parts of the Frailes Formation as
			used by Kaye (1959) and to exclude the younger La Muda Member. Includes Martin González Lava and
			Leprocomio Siltstone Members. Age changed from Late
			Cretaceous(?) to Late Cretaceous, (Pease, 1968a.)
Hackett Sandstone Member (of Hinton For-	Late Mississippian	Virginia and West Virginia	Hackett Sandstone Member of Reger and Price (1926)
mation).	2011 B - 1 1	m	adopted. (Englund, 1968a).  Ivydell Sandstone Member adopted as member of Hance
			Formation (Findland 1968b.)
Hansar Lake Gneiss	Precambrian	Washington and Idaho	Hauser Lake Gneiss adopted. (Weis, 1968.) Hells Half Acre Tuff Member adopted, underlies Apsey
Hells Half Acre Tuff Member (of Galiuro	Miocene	Southeastern Arizona	Hells Half Acre Tuff Member adopted, underlies Apsey
Valorian)			Conglomorate Member (Krieger 1969a)
Henson Formation	Oligocene	Colorado	Age changed from middle and late Tertiary to Oligocene. (Luedke and Burbank, 1968.)
Hildrethe Fermation	Devonian(2)	Maine	Hildreths Formation adopted (Osberg and others, 1968.)
Hinsdale Formation	late Tertiary	Colorado	Hildreths Formation adopted. (Osberg and others, 1968.) Age changed from Pliocene(?) to late Tertiary. (Steven
			and others, 1967.)
Hinton Formation	Late Mississippian	Virginia and West Virginia	Hinton Formation in Virginia and West Virginia divided
			into (in ascending order): Stony Gap Sandstone Member, Hackett Sandstone Member, Little Stone Gap Mem-
			ber, Neal Sandstone Member, middle shale member,
			Tallery Sandstone Member, Pratter Shale Member,
			Falls Mills Sandstone Member, and upper shale mem-
77 1 77 - 1	0	TI-it - 1 Obobon	ber. (Englund, 1968a.)
Holocene Epocn	Quaternary	. United States	Holocene Epoch adopted to replace Recent Epoch as the second and younger epoch in the Quaternary Period.
			(Cohee, 1968.)
Holy Joe Member (of Galiuro Volcanics)	Miocene	Arizona	(Cohee, 1968.)  Holy Joe Member adopted. (Krieger, 1968b.)  Homerian Stage (Floral) adopted. (Wolfe and others, 1966.)  Horseshoe Shale Member adopted. (Mallory, 1967.)
Homerian Stage (Floral)	Miocene and Pliocene(?)	South-central Alaska	- Homerian Stage (Floral) adopted. (Wolfe and others, 1966.)
Horseshoe Shale Member (of Amsden Forma-	Late Mississippian and Early	w yoming	- Horseshoe Shale Member adopted. (Mahory, 1967.)
Huerto Formation	Oligocene	Colorado	Age changed from middle or late Tertiary to Oligocene.
	0-6-0010		(Steven and others, 1967.)  Ijamsville Phyllite is excluded from the Glenarm Series.
Ijamsville Phyllite	early Paleozoic(?)	Maryland	Ijamsville Phyllite is excluded from the Glenarm Series.
Indian Wills Valsanies (of Alder Group)	Procembrian	Arigona	(Southwick and Fisher, 1967.) Indian Hills Volcanics abandoned. Rocks previously
indian ritis voicanies (of Alder Group)	Frecamorian	Al izona	assigned to the Indian Hills Volcanics are now assigned
			to the Green Gulch Volcanics. (Anderson and Creasey,
			1967.)
	Late Mississippian	Virginia_and_West Virginia	Indian Hills Sandstone Member of Reger and Price (1926) adopted. (Englund. 1968a).
Formation).			octobooc. (mukitiid, 1900g).

Name	Age	Location	Revision and reference
Indian Trail Formation	Miocene	Nevada	. Age changed from Miocene and Pliocene(?) to Miocene.
Iron River Iron-Formation Member (of Michigamme Slate).	Precambrian	Michigan and Wisconsin	(Kistler, 1968.)  Iron River Iron-Formation Member abandoned. (James and others, 1968.)
Ivydell Sandstone Member (of Hance Formation).			. Ivydell Sandstone Member adopted. (Englund, 1968b.)
James Run Gneiss  Jawbone Conglomerate Member (of Kiawa Mountain Formation).	late Precambrian(?) Precambrian	. Maryland New Mexico	James Run Gneiss adopted. (Southwick and Fisher, 1967.) Jawbone Conglomerate Member of Barker (1958) adopted. (Barker, this report, p. A21.)
Jeffersonville Limestone			Age changed from Middle Devonian to Early and Middle Devonian (Rouget and Johnson, 1968.)
		•	<ul> <li>Jelm Formation assigned to the Chugwater Group; limited geographically to areas south of the Wind River basin. Members are Red Draw and Sips Creek. (Pipiringos, 1968.)</li> </ul>
Jim Mountain Member (of Wapiti Formation)	Tracero		JimMountainMember adopted. (Nelson and Pierce, 1968.)
Jobos Formation. Josephine Canyon Diorite. Junction Creek Sandstone.	Eocene(?) Late Cretaceous Late Jurassic	Puerto Rico	Jobos Formation adopted. (Nelson, 1967.) Josephine Canyon Diorite adopted. (Drewes, 1968.) Junction Creek Sandstone reduced to member rank as Junction Creek Member of the Wanakah Formation.
Kane Wash Tuff	. Miocene	Nevada	(Hansen, 1968.) Kane Wash Formation of Cook (1965) adopted as Kane
Kanouse Sandstone	Early Devonian	New York and New Jersey	Wash Tuff. (Noble, 1968.)  Age changed from Early and Middle Devonian to Early Devonian. (Boucot and Johnson, 1968.)
Keevy Peak Formation	Precambrian or Paleozoic	Alaska	Vocan Dook Cormetion adopted for rocks that Were
Kiawa Mountain Formation	. Precambrian	. NewMexico	previously part of Birch Creek Schist. (Wahrhaftig, 1968.) Kiawa Mountain Formation of Barker (1958) adopted. (Barker, this report, p. A21.)
Kinnikinic Quartzite	. Middle Ordovician	. Idaho	Age changed from Late Ordovician to Middle Ordovician. (Ruppel, 1998.)
			Kinnikinić Quartzite restricted to uppermost pure well- sorted quartzite of Middle Ordovician age in central Idebo (Hobbs and others 1968)
Kodiak Formation Kokoruda Ranch Complex Kona Dolomite	Cretaceous Tertiary or Cretaceous middle Precambrian	Southern Alaska Western Montana Michigan	Kodiak Formation adopted. (Moore, 1969.) Kokoruda Ranch Complex adopted. (Smedes, 1966.) Kona Dolomite is placed in the Chocolay Group (Animikie
Kummerian Stage (Floral) La Garita Quartz Latite	early Oligocene	West-central Washington	Series). (Gair and Thaden, 1968.) Kummerian Stage (Floral) adopted. (Wolfe, 1968.) Age changed from middle or late Tertlary to Oligocene.
Lake Fork Quartz Latite	Oligocene or older	Colorado	(Steven and others, 1967.).  Name changed to Lake Fork Formation; age changed from Miocene(?) to Oligocene or older. (Olson and others, 1968.)

LaMuda Limestone Member (of Frailes Forma-	Late Cretaceous	Puerto Rico	La Muda Limestone Member raised to formation rank, La Muda Formation. (Pease, 1968a.)
tion).	middle Eocene	Texas	In the Rio Grande embayment south of the Frio River.
Diagono 1 ormanion (or ormanion or orgy, 1			Formation which is restricted to eastern and central
			Charge month of the Frie River (Fordia 1968)
Leprocomio Limestone Member (of Frailes	Late Cretaceous	Puerto Rico	Leprocomio Limestone Member of Frailes Formation revised to Leprocomio Siltstone Member of Guaynabo
Formation).			Formation (Page 1968a)
Lexington Limestone	Middle and Late Ordovician	Kentucky	Age changed from Middle Ordovician to Middle and Late Ordovician. (Cressman, 1967).
Lighthouse Point Member (of Mona Schist)	early Precambrian	Michigan	Lighthouse Point Member adopted. (Gair and Thaden,
THE THE COUNTY	Decembries	South Dakota	1968.) Little Elk Granite of Taylor (1935) adopted. (Zartman and Stern. 1967.)
			and Stern, 1967.) Formation extended into northeastern Utah. (Sando, 1967.)
Little Flat Formation (of Chesterfield Range	Late Mississippian	Northeastern Utah and southeastern Idaho.	
Group).	Early Devonian	New Hampshire	Age of the Littleton, in its type area only, changed from Late Silurian(?) and Early Devonian to Early Devonian.
			The bit is it is a second of the second of t
Loues Formation	Precambrian	South Dakota	Loues Formation adopted. (Redden, 1968.)
Loveland Loess	Pleistocene	Wisconsin and Illinois	Loues Formation adopted. (Redden, 1968.) Loveland Loess changed to Loveland Silt in Illinois and Wisconsin. (Frye and others, 1968.)
Lubbe Creek Formation	Early Jurassic	Alaska	Lubbe Creek Formation adopted. (MacKevett, 1969.)
Lulbegrud Shale Member (of Crab Orchard For-	Middle Silurian	Kentucky	Lubbe Creek Formation adopted. (MacKevett, 1969.) Lubbegrud Shale Member of Foerste (1905) adopted. (Simmons, 1967.)
Lyons Valley Member (of Popo Agie Forma-	Late Triassic	Wyoming.	Lyons Valley Member adopted. (Pipiringos, 1968.)
tion). McClure Mountain Complex	Precambrian or Cambrian	Colorado	McClure Mountain Complex adopted. (Shawe and Parker,
		Y1.1.	Mackey Granite adopted (Nelson and Ross, 1968.)
Mackay Granite	Late Cretaceous	Southeastern Arizona	Madera Canyon Granodiorite adopted (Drewes, 1968.)  Bull Ridge adopted as upper member of the Madison
Madison Limestone	. Early Mississippian (early	Mest-central M Acmme	Timestane (Sanda 1968)
Medrid Formation	Meramec). Silurian or Devonian	Maine	Madrid Formation of Cariani (1959) adopted. (Osberg and
Madrid Formanion	T 1 D 1	Dennamiwania and Ohio	Wherever the Ames Limestone Member (of the Glenshaw
Mahoning Sandstone Member (of Conemaugh Formation).	Late Pennsylvanian	Pennsylvania and Onio	Formation) can be recognized, the Conemaugh is con-
rolmanon,			sidered of group rank and consists of the Casselman and Glenshaw Formations, and the Mahoning is considered
			a member of the Glenshaw Formation. (Roen and others,
		Durate Dies	1968.) Marney Lava Member adopted. (Pease, 1968a.)
Mamey Lava Member (of Camarones Sand-	Late Cretaceous	Puerto Rico	Mamey Lava Member adopted. (Pease, 1968a.)
Mammoth Mountain Rhyolite	Oligocene	. Colorado	Age changed from middle or late Tertiary to Oligocene. (Steven and others, 1967.)
Mancos Shale	Late Cretaceous	New Mexico	Comillo Sandstone Member adopted: overlies Greenhorn
			In Care Trees Dealer (Done and others 1069)
Marhuro Schist	early Paleozoic(?)	_ Maryland and Pennsylvania.	Marburg Schist is excluded from Glenarm Series. (South-
with a my B pattern and a second and a second a		-	wick and Fisher, 1967.)

Name	Age	Location	Revision and reference
Martin González Lava Member (of Guaynabo Formation).	Late Cretaceous	Puerto Rico	Martín González Lava Member adopted. (Pease, 1968a.)
Mayflower Hill Formation.	Early Silurian	_ Maine	. Mayflower Hill Formation adopted. (Osberg and others.
Menominee Group	middle Precambrian	Michigan	1968.) - Ajibik Quartzite and Siamo Slate placed in Menominee
Mesnard Quartzite (of Chocolay Group, Animikie Series).	do	do	Group in the Marquette area. (Gair and Thaden, 1968.)  Mesnard Quartzite placed in Chocolay Group. Former age was Precambrian. Formation restricted to the relatively thick and massive vitreous quartzite. The lower part of the former Mesnard is placed in the Enghantment Leka
Middle Park Formation.	Paleocene	. Colorado	Formation (new). (Gair and Thaden, 1968.)  In the southwestern part of the North Park area, the Middle Park Formation is reduced in rank and made the basal member of the Coalmont Formation. (Hail, 1968.)
Milligen Formation	Early Mississippian	Idaho	Age changed from Devonian(?) and Early Mississippian to
Mingo Formation (of Breathitt Group)	Middle Pennsylvanian	Tennessee	Early Mississippian. (Sandberg and others, 1967.)  Pioneer Sandstone of Glenn (1925) adopted as Pioneer Sand-
Minnelusa Formation			stone Member of Mingo Formation. (Englund, 1968b.)  Age changed from Pennsylvanian and Permian to Early, Middle, and Late Pennsylvanian and Early Permian.
Mona Schist	early Precambrian	Michigan	(Mallory, 1967.)  Mona Schist divided into Lighthouse Point Member and
Monacillo Formation	Late Cretaceous	Puerto Rico	unnamed lower member. (Gair and Thaden, 1968.)  Truiillo Alto Limestone reduced in rank and made member.
Monroe Canyon Limestone	Mississippion	Month contains Their and	of the Monacillo Formation. (Pease, 1968a).
Montgomery Formation	Pleistocene	southeastern Idaho. Texas and Louisiana	Utah. (Sando, 1967.)  Montgomery Formation extended into Texas. (Wilson,
Moose Creek Member (of Totatlanika Schist) Morgan Formation	Daily and Middle	Alaska Wyoming	Moose Creek Member adopted. (Wahrhaftig, 1968.)  Age changed from Middle Pennsylvanian to Early and
$\label{lem:morgantown} \mbox{Mongantown Sandstone Member (of Cone maugh Formation).}$			Middle Pennsylvanian. (Mallory, 1967.)  Wherever the Ames Limestone Member of the Glenshaw Formation can be recognized, the Conemaugh is con- sidered of group rank and consists of the Casselman and Glenshaw Formations, and the Morgantown is con- sidered a member of the Casselman Formation. (Roen
Morton Gneiss	. Precambrian	_ Minnesota	and others, 1968.)  Morton Gneiss of Theil and Dutton (1935) adopted. (Stern
			and others, 1966.)  Morton Loess of Frye and Willman (1960) adopted. (Frye
			and others, 1968.)  Mountain City Formation of Granger and others (1957) adopted. (Coats, 1969.)

Mount Selman Formation (of Claiborne Group)	middle Pesses	m -	
	- middle rocene	. Texas	<ul> <li>Mount Selman Formation abandoned, and former members of formation raised to formation rank. In eastern and</li> </ul>
			central Texas, they include (in ascending order); Reklaw
			Formation, Queen City Sand, and Weches Formation.
			In the Rio Grande embayment south of the Frio River, they include (in ascending order): Bigford Formation
Mount Wrightson Formation	m · · · ·		
			and El Pico Clay (new). (Eargle, 1998.) - Mount Wrightson Formation adopted. (Drewes, 1968.) - Mystic Creek Member adopted. (Wahrhaftig, 1968.)
Naraniita Formation	7.1	and Georgia.	1968.)  Naranjito Formation adopted. (Pease, 1968a.)  Narrow Cape Formation adopted. (Moore, 1969.)
Narrow Cape Formation	Miggene	- Puerto Rico	- Naranjito Formation adopted. (Pease, 1968a.)
Neal Sandstone Member (of Hinton Formation)	Late Mississippian	Virginia and West Virginia	Narrow Cape Formation adopted. (Moore, 1969.) Neal Sandstone Member adopted as substitute for Avis
·	••		Sandstone of Reger and Price (1926). Name Avis preoccu-
Nelson Formation	đo	Namada	pied by reason of use in Texas. (Englund, 1968a.)  Nelson Amphibolite of Granger and others (1957) adopted
		. Nevada	as Nelson Formation. (Coats, 1969.)
Newman Lake Gneiss	Precambrian	Eastern Washington	
146W IVIVEL POPULACION	Early Pennsylvanian	- Pennsylvania, Virginia, and	Pineville Sandstone Member of Hennen and Gawthrop
NizinaMountain Formation	Middle and Late Jurassic	Alaska	Pineville Sandstone Member of Hennen and Gawthrop (1915) adopted. (Englund, 1968a).  Nizina Mountain Formation adopted. (MacKevett, 1969.)
Northeraft Formation	late Eocene	. Washington	Northerait Formation placed in Puget Group. (Gard,
North Haven Greenstone	Middle Ordonicion (2)	Maine	1968.)
37	initiatio oraginatin()	- Mante	(Paylides and others 1968)
Nugget Sandstone	Triassic(?) and Jurassic(?)	_ Wyoming	Age changed from Cambrian(?) to Middle Ordovician(?). (Pavildes and others, 1968.)  Bell Springs Member of Triassic(?) age adopted as basal
			member. Age changed from Early Jurassic to Triassic(?)
Nussbaum Alluvium	Pleistocene	_ Colorado	member. Age changed from Early Jurassic to Triassic(?) and Jurassic(?). (Pipiringos, 1968.)  Age changed from Pleistocene(?) to Pleistocene. (Soister,
O'Brien Creek Formation	middle Teens	37. 41	1967.)
O Dilon Clook Politication	middle Eocene	_ Northeastern Washington	1967.) Age changed from Eocene(?) to middle Eocene. (Yates and Engels. 1968.)
Ohanapecosh Formation	Oligocene	. Washington	and Engels, 1968.)  Age changed from late Eocene to Oligocene. (Wolfe, 1938.)  Ortega Quartzite of Barker (1958) adopted. (Barker, this
Ortega Quartzite	Precambrian	New Mexico	- Ortega Quartzite of Barker (1958) adopted. (Barker, this
Outlet Tunnel Member (of La Garita Quartz	Oligogene	Colorado	report, p. A21.)
Latite).	1 1 25		Age changed from Miocene(?) and Pliocene to late Miocene.  Age changed from Miocene(?) and Pliocene to late Miocene.
Pan Canyon Member (of Paintbrush Tun)	. late Miocene	. Nevada	Age changed from Miocene (?) and Pliocene to late Miocene.
Paintbrush Tuff (of Piapi Canyon Group)	do	do	(Kistler, 1983.)  Age changed from Miocene(?) and Phocene to late Miocene.
Dájaros Tuff	Mandan A. Tat. G. 4		Age changed from Miocene (?) and Pilocene to late Miocene. (Kistler, 1968.) Pájaros Tuff adopted. (Pease, 1968a). Age changed from middle and late Miocene to Pliocene and
Palm Spring Formation	Pliocene and Pleistocene	Southeastern California	Pájaros Tuff adopted. (Pease, 1968a).
D 1 3	1 11000110 1114 1 10101000110	L Doddineastern Camorma	Pleistocene. (Allen and others, 1968.)
Pamunkey Group	Paleocene and Eocene	Maryland, Delaware, and	Brightseat Formation added as basal formation of group:
		Virginia.	age changed from Eocene to Paleocene and Eocene. (Hazel, 1968.)
Pantano Formation	early Oligocene to early	Southeastern Arizona	Pantano Formation of Tolman as defined by Brennan
	Miocana		
Formation).	r ateocette	. waryiand and virginia	Age changed from early Eocene to Paleocene. (Hazel, 1968.)
•			

Name	Age	Location	Revision and reference
Penters Chert	Early Devonian	Arkansas	Age changed from Early or Middle Devonian to Early
			Devonian. (Boucot and Johnson, 1968.)  Peoria Loess of Frye and Willman (1960) adopted. (Frye
Perry Mountain Formation	Silurian(?)	. Maine	and others, (1968).) Perry Mountain Formation of Cariani (1959) adopted.
			(Osberg and others, 1968.)  Peters Creek Quartzite abandoned in Maryland; remains in good usage in Pennsylvania and Virginia. (Southwick and Fisher. 1967.)
Phoenix Park Member (of La Garita Quartz Latite).	Oligocene	Colorado	Age shenged from middle or lote Continue to Oligonome
Piapi Canyon Group	Ditagona		(Steven and others, 1967.)  Age changed from Miocene (?) and Pliocene to late Miocene and early Pliocene. (Kistler, 1968.)
			Age changed from middle and late Tertiary to Oligocene.
			Piña Siltstone Member adopted. (Pease, 1968a).  Lower sandstone beds removed from Redwater Shale  Member and named Pine Butte Member. (Pipiringos,
			Pineville Sandstone Member of Hennen and Gawthrop
			Pioneer Sandstone Member of Glenn (1925) adopted.
Formation)	rateocene	maryland and virginia	Piper Gulch Monzonite adopted. (Drewes, 1968.) Age changed from early Eocene to Paleocene. (Hazel, 1968.)
Pitts Meadow Granodiorite	Precambrian	Colorado	Pitts Meadow Granodiorite adopted. (Hansen, 1968.) Raised in rank to Popo Agie Formation of Chugwater Group. Formation redefined and restricted to analcime-
Poxono Island Formation.	Late Silurian	Northeastern Pennsylvania, northwestern New Jersey, and southeastern New York.	rich beds of Keller (1982). (Pipiringos, 1968.) Poxono Island Shale of White (1882) adopted as Poxono Island Formation. (Epstein and others, 1967.)
ride bhate Member (of Bluestone Formation).	ao	Virginia and West Virginiadodo	Pratter Shale Member adopted. (Englund, 1968a.) Pride Shale Member of Reger and Price (1926) adopted.
Pringle Andesite	Oligocene	Southern Colorado	(Englund, 1968a.) Name changed from Pringle Andesite to Pringle Latite; age changed from Tertiary to Oligocene. (Steven and
Puget Group	Eccene and Oligocene (?)	Washington	In Pierce County the revised Puget Group includes (in ascending order): the Carbonado, Northeraft, and Spiketon Formations. The Wilkeson and Burnett Formations,
Do	early Eocene to early Oligo- cene.	do	formerly in Puget Group, are abandoned. (Gard, 1968.)  In King County the age of the Puget Group is changed from Eccene and Oligocene(?) to (1) early Eccene to

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early Oligocene in the Green River area, and (2) middle Eccene to early Oligocene in the Tiger Mountain area.

raised to formation rank, Queen City Sand; Mount Selman Formation abandoned. (Eargle, 1968.)

(Wolfe, 1968.)

Wyoming \_\_\_\_\_ Age changed from Pennsylvanian to Middle Pennsylvanian

(Des Moines). (Mallory, 1967.)

(Hobbs and others, 1968.)

Wyoming \_\_\_\_\_ Ranchester Limestone Member adopted. (Mallory, 1967.)

ley Formation. (Osberg and others, 1968.)

	•		as second and younger epoch of the Quaternary Period.
Red Draw Member (of Jelm Formation)	Late Triassic	W voming	(Cohee, 1968.) Red Draw Member adopted. (Pipiringos, 1968.)
Red Peak Member (of Chugwater Formation)	Early Triassic	do	Raised to formation rank, Red Peak Formation of Chug-
	Late Jurassic	do	water Group. (Pipiringos, 1998.) Lower sandstone beds removed from Redwater Shale Member and named Pine Butte Member. (Pipiringos,
tion).			1968.)
Reklaw Member (of Mount Selman Formation).	middle Eocene	Texas	Reklaw Member of Mount Selman Formation raised to formation rank, Reklaw Formation; Mount Selman Formation abandoned. (Eargle, 1968.)
Renton Formation (of Puget Group)	late Eocene and early Oligo- cene.	Washington	Age changed from late Eocene and Oligocene(?) to late Eocene and early Oligocene. (Wolfe, 1968.)
Reservation Hill Formation	Pennsylvanian(?) and Permian(?).		Reservation Hill Formation adopted. (Coats, 1969.)
Richland Loess	late Pleistocene		Richland Loess of Frye and Willman (1960) adopted. Age is late Wisconsinan. (Frye and others, 1968.)
			Age changed from Paleocene to Pleistocene. (Steven and others, 1967.)
Río de la Plata Sandstone			Río de la Plata Sandstone of Lidiak (1965) adopted. (Pease, 1968a.)
mation).			Age changed from Late Cretaceous to Early and Late Cretaceous. (Pease, 1968b.)
Rio Piedras Siltstone	Paleocene and Eocene(?)	do	Río Piedras Siltstone adopted. (Pease, 1968a.) Rocket Wash Member adopted. (Noble and others, 1968.)
Root Glacier Formation	Late Jurassic	Southern Alaska	Root Glacier Formation adopted. (MacKevett, 1969.)
Rosita Andesite	Oligocene	Southern Colorado	Name changed from Rosita Andesite to Rosita Formation; age changed from Eocene to Oligocene. (Steven and Epis, 1968.)
Roslyn Formation	middle Eocene	Washington	Age changed from Eocene to middle Eocene. (Wolfe, 1968.)
Roxana Silt	. late Pleistocene	Illinois and Wisconsin	Roxana Silt of Frye and Willman (1960) adopted as basal unit of Wisconsinan Stage. Age is early Wisconsinan (Altonian). (Frye and others, 1968.)

Queen City Sand Member (of Mount Selman middle Eccene. Texas. Queen City Sand Member of Mount Selman Formation

Raging River Formation do... Washington Age changed from middle and late(?) Eccene to middle Eccene. (Wolfe, 1968.)

Rainier Mesa Member (of Timber Mountain early Pliocene. Nevada Age changed from Pliocene to early Pliocene. (Kistler,

Rangeley Formation Silurian(?). Maine Rangeley Conglomerate of Smith (1923) adopted as Range-

Ravenian Stage (Floral) late Eccene West-central Washington Ravenian Stage (Floral) adopted. (Wolfe, 1968.)

Recent Epoch Quaternary United States Recent Epoch abandoned and replaced by Holocene Epoch

Quadrant Quartzite Middle Pennsylvanian (Des

Ranchester Limestone Member (of Amsden Early and Middle Pennsyl-

Formation).

Formation).

Tuff).

Moines).

vanian.

Name	Age	Location	Revision and reference
			Salero Formation adopted. (Drewes, 1968.)  Age changed from Early or Middle Devonian to Early
Formation).			Devonian. (Boucot and Johnson, 1968.) Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemaugh is con- sidered of group rank and consists of the Casselman and Glenshaw Formations, and the Saltsburg is considered a nnember of the Glenshaw Formation. (Roen and others, 1968.)
			Age changed from middle or late Tertiary to Oligocene or
			Age changed from Eocene(?) to middle Eocene. (Yates and
			Age changed from Oligocene to middle Oligocene. (Briggs,
Santa Ólaya Lava	Late Cretaceous Oligocene Middle Ordovician and	Puerto Rico Colorado Idaho	Santa Olaya Lava of Lidiak (1965) adopted. (Pease, 1968a.) Sapinero Mesa Tuff adopted. (Olson and others, 1968.) Age changed from Late Ordovician to Middle Ordovician
			and younger. (Ruppel, 1968.)  Name changed to Schoharie Formation. (Boucot and
Schuler Formation (of Cotton Valley Group)	Late Jurassic	Texas, Louisiana, and Arkansas.	Johnson, 1998.)  Schuler Formation redefined to include all rocks between the top of the Buckner Formation and the base of the Cretaceous rocks, except for the Q tongue of the Bossier
Seldovian Stage (Floral) Semilla Sandstone Member (of Mancos Shale)	Oligocene(?) and Miocene Late Cretaceous	South-central Alaska New Mexico	Seldovian Stage (Floral) adopted. (Wolfe and others, 1966.) Semilla Sandstone Member adopted. (Dane and others,
Setters Formation (of Glenarm Series)	late Precambrian(?)	Maryland, Pennsylvania, and	1968.) Age changed from early Paleozoic(?) to late Precambrian(?).
Shallow Creek Quartz Latite	Oligocene	Delaware. Colorado	(Southwick and Fisher, 1967.) Age changed from middle or late Tertiary to Oligocene.
Sheep Creek Member (of Totatlanika Schist) Shellenberger Canyon Formation (of Bisbee Group).	Mississippian(?) Early Cretaceous	Central AlaskaSoutheastern Arizona	(Steven and others, 1967.) Sheep Creek Member adopted. (Wahrhaftig, 1968.) Shellenberger Canyon Formation of Tyrrell (1957) adopted. (Finall this result of A22.)
Sheppard Granite	Eocene	Northeastern Washington	(Finnell, this report, p. A32.) Age changed from Tertiary to Eocene. (Yates and Engels,
Siamo Slate	middle Precambrian	Michigan	1968.) Placed in Menominee Group of Animikie Series. Age changed from Precambrian to middle Precambrian.
Siletz River Volcanic Series	early and middle Eocene	Western Oregon	(Gair and Thaden, 1968.) Name changed from Siletz River Volcanic Series to Siletz
Silver Run Limestone	early Paleozoic(?)	Maryland	River Volcanics. (Snavely and others, 1968.) Silver Run Limestone is excluded from Glenarm Series.
Silverton Volcanic Group	Oligocene	Colorado	(Southwick and Fisher, 1967.) Age changed from middle and late Tertiary to Oligocene. (Luedke and Burbank, 1968.)

			and the second second
Sine Creek Member (of Jelm Formation)	Late Triassic	. Wyoming	Sips Creek Member adopted. (Pipiringos, 1968.) Sitkelidak Formation adopted. (Moore, 1969.)
Sitkalidak Formation	Oligocene	do	Sitkinak Formation adopted. (Moore, 1909.)
Smalls Falls Formation	Shuran(:)	W13111C	1 -11 -1 - 1000
		a. 1. 1.	Age shound from middle or lete Tertiary to Oligogene
Snowshoe Mountain Quartz Latite	Oligocene	. Colorado	Age changed from middle or late Tertiary to Oligocene. (Steven and others, 1967.)
		Wassing and Montons	Grove Creek Formation reduced to member rank and made
Snowy Range Formation	Late Camprian	. Wyoming and Montana	upper member of Snowy Range Formation. (Pierce and
			NTalean 1968 )
	Tata(9) Chatagoons	Northwestern California	South Fork Mountain Schist adopted. (Blake and others,
South Fork Mountain Schist	Date(I) Cletaceous	. Itorra description	1967.)
Spiketon Formation (of Puget Group)	lata(2) Eocene	Washington	Spiketon Formation adopted. (Gard, 1968.)
Spiketon Formation (of Fuget Group)	Turassic	Southeastern Arizona	Squaw Gulch Granite adopted. (Drewes, 1968.)
Star Park Formation	Early to Late Triassic.	Nevada	
Dual I Cak Pullbaulon			Late Triassic. (Silberling, 1968.)
Steens Basalt	middle(?) and late Miocene	. Oregon	Age changed from middle Pliocene to middle(?) and late Miocene. (Walker and Swanson, 1988.)
DUOOID DUOOLU LLLLLLLLLLLLLLLLLLLLLLLLLLLLLLL			Miocene. (Walker and Swanson, 1968.)
Stockade Wash Member (of Paintbrush Tuff)	. late Miocene	. Nevada	Age changed from Miocene(?) and Pliocene to late Miocene. (Kistler, 1968.)
		Gtht A sisono	Ago changed from Cretacours or Testiary to Late Creta-
Sugarloaf Quartz Latite	. Late Cretaceous	Southeastern Arizona	Age changed from Cretaceous or Tertiary to Late Cretaceous, (Hayes and Drewes, 1968.)
	Middle Oudemision	Control Ventucky	Sulphur Well Member of McFarlan (1943) adopted. (Cress-
Sulphur Well Member (of Lexington Limestone).	. Middle Ordovician	Central Kentucky	man, 1968.)
a 1 22 a deles Mentes (of Consmonath	Tata Danneylvanian	Pennsylvania and Ohio	Wherever the Ames Limestone Member (of the Glenshaw
Summernin Sandstone Member (of Cohemadgi	Late I emisyivaman	_ 1 onnog1, units units =	Formation) can be recognized, the Conemaugh is con-
Formation).			sidered of group rank and consists of the Casselman and
			Glenshaw Formations, and the Summerhill is considered
			a member of the Casselman Formation. (Roen and
			others, 1968.)
Sundance Formation	Late Jurassic	_ Wyoming	Two new members adopted—Pine Butte between Lak and Redwater and Windy Hill at top of formation. (Pipi-
Duradiio 2 dimensional			
			ringos, 1968.)
Sykesville Formation	_ Paleozoic	_ Pennsylvania	Sykesville Formation abandoned in Maryland. Rocks are now included in boulder gneiss lithofacies of Wissahickon
·			Formation, Sykesyille Formation still in good usage in
			Poppsylvania (Southwick and Fisher 1967.)
	T -4-3 final almoston	Virginia and Wast Virginia	Tallery Sandstone Member of Reger and Price (1926)
Tallery Sandstone Member (of Hinton Forma-	Late Wississippian	. Vilgilla and West Vilgilla.	adopted. (Englund, 1968a.)
tion).	Forly Parmian (Lagnard)	Teras	Age changed from Early Permian (Leonard?) to Early Permian (Leonard). (Myers, 1968.)
Tarpa Limestone Member (of Clyde Formation)	- Early 1 elillair (Deonard)	_ 104.00	Permian (Leonard). (Myers, 1968.)
Dellawide Formation	early Eocene	Colorado	Age changed from Oligocene(?) to early Eocene. (Steven and others, 1967.)
1 enuriue Formation	_ 0413 1300020111111111111111111111111111111		and others, 1967.)
Temporal Formation	Early Cretaceous	Southeastern Arizona	and others, 1967.)  Temporal Formation adopted. (Drewes, 1968.)  Age changed from middle (?) and late Eocene to middle and late Eocene. (Wolfe, 1968.)
Tiger Mountain Formation (of Puget Group)	_ middle and late Eocene	Washington	Age changed from middle(?) and late Eocene to middle and
Tiget intomitation (or a second to the first			late Eocene. (Wolfe, 1968.)
Timber Mountain Tuff (of Piapi Canyor	early Pliocene	. Nevada	Age changed from Pliocene to early Pliocene. (Kistler, 1968.)
Group).		a	Age changed from Microne(?) and Plicene to late Microne.
Tiva CanyonMember (of Paintbrush Tuff)	late Milocene	u0	. Age changed from Miocene(?) and Pliocene to late Miocene. (Kistler, 1968.)
	Mantiana an Orratamana	Southeastern Alaska	. Tlevak Basalt adopted. (Eberlein and Churkin, this
Tlevak Basalt	_ rerusty or quaternary	_ DUUDIIGASTOITI ATASKA	report, p. A25.)
			EDEC OF A ENERGY

Name	Age	Location	Revision and reference
Topopah Spring Member (of Paintbrush Tuff)	. lateMiocene	Nevada	Age changed from Miocene (?) and Pliocene to late Miocene.
			(Kistler, 1968.) Stratigraphically restricted to exclude rocks assigned to Keevy Peak Formation. Subdivided into five newly named members (in ascending order): Moose Creek, California Creek, Chute Creek, Mystic Creek, and Sheep Creek. (Wahrhaftig, 1968.)
			Age changed from middle or late Tertiary to Oligocene. (Steven and others, 1967.)
Trident Member (of Three Forks Formation)			Extended to Idaho. (Sandberg and others, 1967.)
		. Wyoming	Trout Peak Trachyandesite adopted. (Nelson and Pierce, 1968.)
Trujillo Alto Limestone	Late Cretaceous	. Puerto Rico	. Trujillo Alto Limestone reduced in rank and made member
Tub Spring Member (of Belted Range Tuff)	late Miocene	Nevada	
Tub Spring Member (of Indian Trail Formation).	Miocene	do	Age changed from Miocene and Phocene(7) to late Miocene. (Noble and others, 1968.)  Age changed from Miocene and Phocene(?) to Miocene.
Tugidak Formation Tukwila Formation	Pliocene late Eocene and early Oligocene.	South-central Alaska Washington	(Kistler, 1968.) Tugidak Formation adopted. (Moore, 1969.) Age changed from late Eocene to late Eocene and early
Turney Ranch Formation (of Bisbee Group)	Early Cretaceous	Southeastern Arizona	Oligocene. (Wolle, 1968.) Turney Ranch Formation of Tyrrell (1957) adopted. Top formation of Bisbee Group. (Finnell, this report,
Twocreekan Substage	late Pleistocene	Illinois and Wisconsin	<ul> <li>Twocreekan Substage of Frye and Willman (1960) adopted as second youngest substage of Wisconsinan Stage.</li> <li>Overlies Woodfordian Substage and underlies Valderan</li> </ul>
Uffington Shale Member (of Conemaugh Formation).	Late Pennsylvanian	Pennsylvania and Ohio	Substage. (Frye and others, 1968.) Wherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemangh is con- sidered of group rank and consists of the Casselman and Glenshaw Formations, and the Uffington is considered a member of the Glenshaw Formation. (Roen and
			others, 1968.)  Age changed from Middle and Late Ordovician to Middle
Urbana Phyllite	early Paleozoic(?)	Maryland	Ordovician. (Bachman, 1968.) Urbana Phyllite excluded from Glenarm Series. (South-
Uyak Formation Valderan Substage	Triassiclate Pleistocene	South-central Alaska Illinois and Wisconsin	wick and Fisher, 1967.)  Uyak Formation adopted. (Moore, 1969.)  Valderan Substage of Frye and Willman (1960) adopted as uppermost substage of Wisconsinan Stage. (Frye and
Vanderlehr Formation	Precambrian	South Dakota	others, 1968.)  Vanderlehr Formation adopted. (Redden, 1968.)

ge changed from early Pliocene to late Miocene to early
Pliocene. (Swain, 1968.) Vaco Limestone of Foerste (1905) adopted. (Simmons
1967.) ge changed from Miocene and Pliocene to late Miocene.
(Kistler, 1968.) Vakefield Marble excluded from Glenarm Series. (South-
wick and Fisher, 1967.) ge changed from Triassic or Jurassic to Late Triassic.
(Haves and Drewes, 1968.)
nnction Creek Sandstone reduced in rank to Junction Creek Member of Wanakah Formation. (Hansen, 1968.) apiti Formation adopted. (Nelson and Pierce, 1968.)
ge changed from middle or late Tertiary to Oligocene. (Steven and others, 1967.)
adopted as Waterville Formation. (Osberg and others,
1968.)
Jebb Formation adopted. (Smith and Ketner, 1968.) ge changed from Pennsylvanian and Permian to Middle and Late Pennsylvanian and Permian. (Mallory, 1967.) Jeches Greensand Member (of Mount Selman Formation) raised in rank to Weches Formation; Mount Selman abandoned. (Eargle, 1968.)
Vedron Formation of Frye and Willman (1960) adopted.  Age is late Wisconsinan (Woodfordian). (Frye and others,
1968.) Vehutty Formation adopted. (Hernon, 1969.)
ge changed from Miocene(?) to Oligocene or older. (Olson
and others, 1968.)  Ige changed from Pennsylvanian or younger to Late
Pennsylvanian or younger. (Feininger, 1968.) Yewe Slate placed in Chocolay Group, Animikie Series.
Age changed from Precambrian to middle Precambrian. (Gair and Thaden, 1968.)
Vilkeson Formation abandoned; strata assigned to the Carbonado Formation (revised), (Gard, 1968.)
ge changed from Late(?) Cretaceous or Tertiary to Late Cretaceous and (or) early Tertiary. (Krieger, 1968b.)
Villow Canyon Formation of Tyrrell (1957) adopted. (Finnell, this report, p. A31.)
Age changed from middle or late Tertiary to Oligocene. (Steven and others, 1967.)
Vherever the Ames Limestone Member (of the Glenshaw Formation) can be recognized, the Conemaugh is consid-
ered of group rank and consists of the Casselman and Glenshaw Formations, and the Wilmore is considered a member of the Casselman Formation. (Roen and
others, 1968.)

Wacca	maw Formation	late Miocene to early Pliocene.	North Carolina and South Carolina.	Age ch
Wacol	Member (of Crab Orchard Formation)	Middle Silurian	Kentucky	Wacol
				1967.
Wahm	onie Formation	late Miocene	Nevada	Age cn (Kist
	eld Marble			Wakefi wick
	t Gap Volcanics			(Hav
Wanak	calı Formation	Late Jurassic	Colorado	Junctio
Wapit	Formation	early(?) and middle(?) Eocene.	Northwestern Wyoming	Wapiti
Wason	Park Rhyolite	Oligocene	Colorado	Age ch
Water	ville Formation	Silurian(?)	Maine	
		.,		adop 1968.
Webb	Formation	Early Mississippian	Nevada	Webb
	Sandstone	nion and Downion		and
Weche For	s Greensand Member (of Mount Selman nation).	middle Eocene	Eastern and central Texas	Weches raised abar
Wedro	n Formation	late Pleistocene	Illinois and Wisconsin	Wedro
				Age 1 1968.
Wehut	ty Formation	Precambrian	and Georgia.	Wehut
West	Elk Breccia	Oligocene or older	Colorado	Age ch
	ly Granite	**************************************	Teland	Age c
Wewe	Slate	middle Precambrian	Michigan	Wewe
				ALGO !
Wilke	son Formation (of Puget Group)	Eocene	. Washington	<b>(Gai</b> Wilkes Carb
Willia	mson Canyon Volcanics	Late Cretaceous and (or) early Tertiary.	Arizona	
Willow	Canyon Formation (of Bisbee Group)	Early Cretaceous	Southeastern Arizona	Willow (Fin:
Willow	7 Creek Member (of Bachelor Mountain volite).	Oligocene	Colorado	Age cl
Wilmo	onter. ree Sandstone Member (of Conemaugh mation).	Late Pennsylvanian	Pennsylvania and Ohio	

Name	Age	Location	Revision and reference
Windy Gulch Member (of Bachelor Mountain Rhyolite).	Oligocene	Colorado	Age changed from middle or late Tertiary to Oligocene. (Steven and others, 1967.)
Windy Hill Sandstone Member (of Sundance Formation).	Late Jurassic	Wyoming	Windy Hill Sandstone Member adopted. (Pipiringos, 1968.)
Winnebago Drift	. late Pleistocene	Northern Illinois	
			is Wisconsinan (Altonian). (Frye and others, 1968.) Wisconsinan Stage of Frye and Willman (1960) adopted. Includes (in ascending order): Altonian, Farmdalian, Woodfordian, Twocreekan, and Valderan Substages. (Frye and others, 1968.)
Wissahickon Formation (of Glenarm Series)	. late Precambrian(?)	Pennsylvania, New Jersey, Delaware, Maryland, and Virginia.	Age changed from early Paleozoic(?) to late Precambrian(?); in Maryland includes strata formerly assigned to Peters Creek and Sykesville Formations which have been abandoned there. (Southwick and Fisher, 1967.)
Woodfordian Substage	late Pleistocene	Illinois and Wisconsin	Woodfordian Substage of Frye and Willman (1960) adopted. Overlies Farmdalian Substage and underlies Twocreekan Substage. (Frye and others, 1968.)
Woodruff Formation Yakima Basalt	Early to Late Devonian late Miocene and early Pliocene.	Nevada Oregon	Woodruff Formation adopted. (Smith and Ketner, 1968.) Yakima Basalt extended into north-central Oregon. (Waters
Yucca Mountain Member (of Paintbrush Tuff)	late Miocene		(Figther 1969)
Zooks Corner Formation	Cambrian	. Pennsylvania	Zooks Corner Formation adopted. (Meisler and Becher, 1968.)

# ORTEGA QUARTZITE AND THE BIG ROCK AND JAWBONE CONGLOMERATE MEMBERS OF THE KIAWA MOUNTAIN FORMATION, TUSAS MOUNTAINS, NEW MEXICO

#### By FRED BARKER

The Tusas Mountains of northern New Mexico contain extensive exposures of Precambrian quartzite and conglomerate. In a reconnaissance investigation Just (1937, p. 43, and pl. III) applied the name Ortega Quartzite to these rocks for their exposures from Ojo Caliente to Jawbone Mountain (T. 29 N., R. 6 E.) with the type locality in the Ortega Mountains. In a more recent study in the Las Tablas, Canon Plaza, Burned Mountain, and Mule Canyon 7½-minute quadrangles, Barker (1954, 1958) subdivided Just's Ortega Quartzite. In conformance with this work, the name Ortega Quartzite is applied only to the 14,000- to 20,000-foot-thick, lower part of this very thick section of quartite and conglomerate as exposed in the center of section 25, T. 27 N., R. 8 E., and to the west and southwest. The base of the Ortega Quartzite is not exposed in the Tusas Mountains. The strata overlying the redefined Ortega Quartzite are named the Kiawa Mountain Formation, after the exposures at Kiawa Mountain, the type locality (secs. 3-5, 8-10, T. 27 N., R. 6 E.; Barker, 1958). The uppermost part of the Kiawa Mountain Formation is not preserved in the Tusas Mountains.

The lowest unit of the Kiawa Mountain Formation in the Las Tablas quadrangle is a pebble conglomerate and is named the Big Rock Conglomerate Member (Barker, 1958) for exposures about 0.2 to 1 mile north and northwest of Big Rock. The type locality is in the SE½ sec. 27, T. 27 N., R. 8 E. About 15 miles to the northwest, in the Burned Mountain quadrangle, the basal part of the Kiawa Mountain Formation is pebble conglomerate, fine-grained conglomerate, and quartzite and is named the Jawbone Conglomerate Member (Barker, 1958). The type locality is in the N½ sec. 19, T. 29 N., R. 7 E. Excellent exposures of this member are found at Jawbone Mountain which lies in the northwest corner of the Burned Mountain quadrangle and in the northeast corner of the Cebolla quadrangle. The stratigraphic relations of these two conglomerates are not known; however, there is a possibility that the Jawbone is younger than the Big Rock (Barker, 1958, pl. 1).

The Ortega Quartzite, restricted, and the Kiawa Mountain Formation and its two members, the Big Rock and Jawbone Conglomerate Members, are herein adopted.

The Kiawa Mountain Formation consists principally of quartzite and contains some units of conglomerate and amphibolite. The quartzite layers are light blue to gray, dense, vitreous, and typically very fine grained to medium grained. Layers containing granules and

pebbles of quartz and, less commonly, of ferruginous chert or ironformation occur throughout. At Kiawa Mountain and to the northwest, the quartzite consists of 90 to 98 percent quartz with minor amounts of hematite, kyanite, manganian and alusite, rutile, and other minerals. Northeast of Kiawa Mountain, in the canyon of Spring Creek, the quartzite is feldspathic and muscovitic. South of Kiawa Mountain, on La Jarita Mesa, the quartzite of this formation grades into the Petaca Schist of Just (1937, p. 43), which is a muscovitic quartzose schist. The Petaca Schist is only found adjacent to the pegmatite bodies of the Petaca District, and Just ascribes the muscovite in the schist to emanations from the pegmatites.

The quartzite layers of the Kiawa Mountain Formation are intensively folded, and so their thickness cannot be measured accurately. However, they are estimated to be 5,000 to 10,000 feet thick (Barker, 1958, p. 10).

The Big Rock Conglomerate Member consists of interlayered gray quartzose conglomerate and pebbly quartzite. The clasts are of quartz and jasper, and they generally range from ½ to 5 inches in maximum dimension. The matrix of both rock types typically is ½ to 4 mm (millimeters) in grain size and shows a mosaic fabric. Quartz forms 60 to 80 percent; microcline and muscovite are next in abundance; and biotite, garnet, and hematite are accessory in the matrix. Both current bedding and cross-stratification are common in the Big Rock Conglomerate Member. This member is intricately folded, and is estimated to be 100 to 200 feet thick near Big Rock. It pinches out to the southeast. To the northwest it also apparently pinches out; here, however, this conglomerate was intruded by Precambrian rhyolite and later was partly covered by Tertiary conglomerate (Barker, 1958, pl. 1), so its stratigraphic relations are obscured.

The Jawbone Conglomerate Member is comprised of interlayered, gray to light gray, vitreous, massive, quartzose conglomerate, quartzite, and pebbly conglomerate. Granules and well-rounded pebbles of quartz and jasper of ½ to 1 inch in grain size form the coarsest part of the conglomerates. The matrix of the conglomerate and the quartzite consists of recrystallized quartz, mostly ½ to 1 mm in grain size, and accessory kyanite, hematite, rutile, and muscovite. Bedding and cross-stratification may be seen in many outcrops of the Jawbone. Because this member is folded, its thickness may not be directly measured; however, it probably is at least 500 feet thick and may be more than 2,000 feet thick.

# REASONS FOR ABANDONMENT OF THE PORTAGE GROUP By Wallace de Witt, Jr.

In our study of the stratigraphy of the Upper Devonian rocks in western and west-central New York (Pepper, de Witt, and Colton, 1956; Colton and de Witt, 1958; de Witt and Colton, 1959; de Witt, 1960), we were more or less continuously entangled with the old "classic" nomenclature which was introduced in the fourth geological district by James Hall (1839, p. 322) and in the third geological district by Lardner Vanuxem (1842, p. 172). Not long after we began to study the rocks of the West Falls Formation—the rocks above the Sonvea Formation and below the Java Formation (Pepper, de Witt, and Colton, 1956)—we realized that the many usages of the name Portage as proposed by geologists during the past 125 years and the correlation of the units involved with the name were a source of confusion and difficulty. At different times, the name Portage had been applied to a single lithologic unit (Luther, 1897), to a group of several lithologic units (Hall, 1840, p. 391; 1843, p. 224), to a series of rocks in a time sense (Tilton and others, 1927, p. 88-90), and in parts of Kentucky. New York, Pennsylvania, and West Virginia to a magnafacies that included most of the marine Upper Devonian rocks. The name was given not only to a group of rocks exposed along the Genesee River in western New York but also to the upper massive sandstone unit within that group (Hall, 1843).

The relationship of the stratigraphic units that make up the Portage Group of Hall (1839) along the Genesee River in the fourth geological district and the units in the Portage or Nunda Group of Vanuxem (1842), about 75 miles to the east near Ithaca in the third geological district, illustrates some of the confusion surrounding the name Portage. Hall and Vanuxem believed the two groups were correlative.

In the type area along the Genesee River, Hall's Portage Group (1839) consisted of the strata above the Gardeau or Lower Fucoidal Group and below the Chemung Group. The Portage consisted of a thin basal sandstone, the Table Rock Sandstone of Chadwick (1933), a medial unit of intercalated shale and sandstone, and a thick upper massive sandstone. These units make up the upper part of our West Falls Formation at the Genesee. The upper massive sandstone is the Nunda Sandstone Member of our West Falls, and the two lower units make up the West Hill Member and the upper part of the Gardeau Shale Member of the West Falls.

In 1842 Vanuxem defined his Portage or Nunda Group near Ithaca as including the Sherburne Flagstone and Shale, the Cashaqua Shale, and the Gardeau and Portage Groups of the previous State reports. Vanuxem drew the top of his Portage Group at the base of the overlying Ithaca Group. Our study of these Upper Devonian rocks showed

that in the vicinity of Ithaca, the strata in the basal part of Vanuxem's Ithaca Group are the Renwick Shale Member, Ithaca Member, and West River Shale Member of our Genesee Formation. These strata are equivalent at the Genesee River to rocks which comprise the upper half of our Genesee Formation—the upper part of Hall's Genesee Slate. Thus, the strata near Ithaca that Vanuxem assigned to the Portage Group are not the lateral equivalents of Hall's Portage on the Genesee but are separated from the Portage beds by the lower part of the West Falls Formation and all the subjacent Sonyea Formation. The strata that compose the upper part of the West Falls Formation on the Genesee, Hall's Portage Group, change facies to the east; and if present in the Ithaca area, the equivalents of Hall's Portage Group occur in the upper part of the Chemung rocks well south and southwest of Ithaca. In their respective districts, both Hall and Vanuxem defined the top of their Portage Group largely by paleontologic criteria. They were not aware that because of changes in lithology the equivalents of Hall's original Portage Group lay within a dissimilar facies with a different fauna in the Ithaca area. If the original definition of the group (Hall, 1839) is strictly followed, none of Vanuxem's Portage at Ithaca is Portage because all the strata are older than the beds originally assigned to the group by Hall.

In 1843 Hall added to the problem by expanding his Portage Group on the Genesee to include the following units in descending order: Gardeau or Lower Fucoidal Group, Cashaqua Shale, and Sherburne Flagstones. The Sherburne Flagstones, the Sherburne Flagstone Member of our Genesee Formation, feathers out of the sequence to the east of the Genesee River in the area between Seneca and Cayuga Lakes and is unrepresented by coarse-grained rocks on the Genesee. As redefined by Hall, his expanded Portage Group includes the upper half of our Genesee Formation and all the younger Sonyea and West Falls Formations.

In the years between 1843 and 1943, the use of Portage Group, Portage Formation, or Portage Series, spread widely across the Appalachian basin as attempts were made by many geologists to synthesize the Upper Devonian stratigraphy of this area. Correlations were proposed partly on the basis of similar lithology and stratigraphic position but more commonly on the presence of similarity of faunas. Usage of the name Portage became so varied that each author was more or less obliged to redefine the unit in each paper. The wide latitude of usage of Portage increased the misunderstanding surrounding the unit. A glance through the papers of the New York Geological Survey and particularly the writing of G. H. Chadwick will illustrate many of the changes in the usage of Portage during the past 125 years.

From these and much other data, we concluded that the name Portage was largely invalid because of the multitudinous usages in the past. Furthermore, within the framework of stratigraphic nomenclature which we have proposed for the Upper Devonian rocks in the western half of New York, Portage need not be recognized in a formal sense. Consequently, we abandoned use of Portage rather than continue to compound the confusion surrounding the name by attempting to rigorously redefine the stratigraphic unit. The limitation on the length of text accompanying oil and gas charts (Pepper and others, 1956; Colton and de Witt, 1958) precluded a lengthy discussion of the stratigraphic nomenclature and possible revision; also, some of the most cogent reasons in favor of abandoning Portage as a group or formation were not clearly evident until we had completed work on the Sonyea Formation and the Genesee Formation.

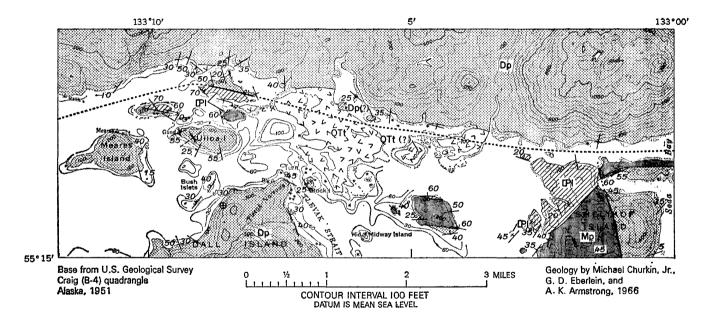
The New York Geological Survey did not include the Portage among the stratigraphic units recognized in the Upper Devonian sequence in the text of the most recent issue of the State geologic map (New York State Museum and Science Service Geological Survey, 1962). The New York Survey included the rocks that made up Hall's expanded Portage Group (Hall, 1843) as parts of our Genesee, Sonyea, and West Falls Formations. On the map, our formations were given group status. Clearly the action of the New York Geological Survey indicates the intent to abandon the Portage. Similarly Rickard's correlation chart of the Devonian rocks in New York (Rickard, 1964) does not contain the Portage as a recognized stratigraphic unit. Rickard shows in the chart that the "Chagrin" ("Portage") is a depositional phase—a gray shale, siltstone, limestone, and calcareous nodule facies—which is present at different places in several of the groups of strata in the Upper Devonian sequence in western New York.

# TLEVAK BASALT, WEST COAST OF PRINCE OF WALES ISLAND, SOUTHEASTERN ALASKA

By G. Donald Eberlein and Michael Churkin, Jr.

A number of small islands off the west coast of Prince of Wales Island at the latitude of Ketchikan (fig. 1) are underlain by one or more basalt flows. These flows are best exposed along the shoreline. The flows make up a new formation, here named the Tlevak Basalt for its occurrence in the vicinity of Tlevak Strait (fig. 1). The basalt is probably flat lying, and the low relief of the shoreline exposes only a few feet of lava at any locality.

Generally the basalt is strongly jointed and forms thin subhorizontal slabs. In places there are well-developed closely spaced concentric fractures along which the rock weathers spheroidally. Locally, the



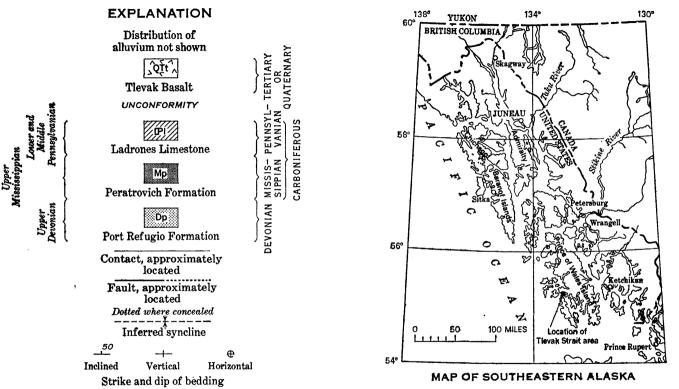


FIGURE 1.—The Tlevak Strait area, southeastern Alaska.

basalt has vertical columnar joints that are transected by the subhorizontal fracture system.

In contrast to the Paleozoic formations which predominate in this region, several of which contain basaltic volcanic rocks, the Tlevak Basalt is remarkably fresh and contains phenocrysts of olivine and plagioclase (labradorite) in an intergranular groundmass of labradorite microlites, olivine, and subordinate clinopyroxene. The rock has a normative mode and may be classified as olivine basalt according to the scheme of Yoder and Tilley (1962).

On the south shore of a small unnamed island about 1 mile northwest of the north entrance to Tlevak Strait, the Tlevak Basalt unconformably overlies steeply dipping beds of the Ladrones Limestone. Elsewhere, in discontinuous exposures, the Tlevak apparently overlies the Port Refugio and Peratrovich Formations. The Tlevak Basalt thus appears to rest unconformably upon Pennsylvanian, Mississippian, and Upper Devonian formations. A small area of similarly fresh olivine basalt is exposed along the northeast shore of Trocadero Bay about 9 miles northeast of the area of this report. This olivine basalt may overlie a formation older than the Port Refugio, the Descon Formation.

The top of the Tlevak Basalt is not exposed, and its total thickness is uncertain. Judging from its inferred gross subhorizontal structure and the maximum relief adjacent to shoreline exposures, it is probably less than 100 feet thick.

The Tlevak Basalt is demonstrably younger than the Ladrones Limestone and therefore is post-Pennsylvanian. It is tentatively assigned a Tertiary or Quaternary age because of its freshness and general lack of deformation.

# FORMATIONS OF THE BISBEE GROUP, EMPIRE MOUNTAINS QUADRANGLE, PIMA COUNTY, ARIZONA

By Tommy L. Finnell

A thick sequence of shale, sandstone, conglomerate, and limestone of Early Cretaceous age is widespread in southeastern Arizona. Outcrops of these sedimentary rocks in the Mule Mountains near Bisbee (fig. 2) were called the Bisbee beds by Dumble (1902, p. 706). Later, they were designated the Bisbee Group by Ransome (1904, p. 56), who distinguished the Glance Conglomerate at the base, the Morita Formation, the Mural Limestone, and the Cintura Formation at the top. However, in areas north of the Mule Mountains, such as the Tombstone Hills and the Dragoon Mountains (Gilluly, 1956, p. 74), and to the northwest in the Mustang Mountains (Hayes and Raup, 1968), no lithologic equivalent of the Mural Limestone occurs, and the Bisbee was mapped as a formation with a basal conglomerate

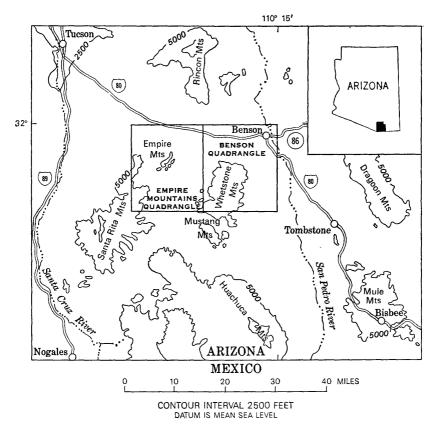


FIGURE 2.—Empire Mountains area, Arizona.

member. Cretaceous strata in the Whetstone Mountains, called Bisbee Group by Tyrrell (1957; 1965) and Bisbee(?) Formation by Creasey (1967), are partly correlative with the Bisbee Group in its type area (Hayes and Drewes, 1968, p. 55); but, in the absence of an equivalent to the Mural Limestone, Tyrrell divided the group into the Willow Canyon Formation at the base, the Apache Canyon Formation, the Shelleburg Canyon Formation, and the Turney Ranch Formation at the top. Schafroth (1968, p. 60-64) recognized the same formations in the Empire and northern Santa Rita Mountains. Recent detailed studies confirm that the formations recognized by Tyrrell are also valid units in the Empire Mountains quadrangle, and the names are therefore adopted for use as described below. However, both earlier workers included the Glance Conglomerate as a member of the Willow Canyon Formation, whereas I have chosen to map the Glance as a separate formation. Hence, the Willow Canyon Formation in the Empire Mountains quadrangle is restricted to

those beds that occur stratigraphically above the Glance Conglomerate and below the Apache Canyon Formation. Also, the Shelleburg Canyon Formation of Tyrrell (1957) is here called Shellenberger Canyon Formation so that the spelling will be the same as that shown on the 1958 edition of the U.S. Geological Survey topographic map of the Benson quadrangle.

#### GLANCE CONGLOMERATE

The Glance Conglomerate, the basal formation of the Bisbee Group, consists of as much as 5,600 feet of pebble to boulder conglomerate and sedimentary breccia of variable lithology. In general, two types of conglomerate occur in the Glance. One is composed mainly of limestone and dolomite fragments and is present along the south, east, and north flanks of the Empire Mountains; the other is composed mainly of quartzite and granitoid fragments and is only present in the northern part of the range where it conformably overlies the limestone conglomerate.

The limestone conglomerate ranges in thickness from 1 to 400 feet and consists of closely packed subangular to subrounded fragments of limestone and dolomite as much as 3 feet across set in a matrix of reddish-brown to gray calcareous sandstone and sandy siltstone. In places, tabular blocks of Paleozoic formations as much as 1,000 feet across are embedded in the conglomerate.

The quartzite-granitoid conglomerate is as much as 5,200 feet thick, and it consists of closely packed angular to subangular fragments of quartzite and gneissic quartz diorite as much as 3 feet across set in a matrix of reddish-brown to greenish-gray sandstone and sandy siltstone. The quartzite fragments are most numerous at the base, giving way gradually upward to the quartz diorite fragments.

The Glance rests on an erosional unconformity of widely variable relief. Along the south and east flanks of the Empire Mountains, the surface of unconformity has irregularities measured in tens of feet that are reflected in the variations of thickness of the conglomerate. In the northern part of the range, however, the unconformity was cut over a mass of Precambrian granitoid and metamorphic rocks that stood several thousand feet higher than the floor of the basin of Bisbee deposition. There, the Glance consists of debris eroded from the highland and deposited along the basin margin. This Glance grades southward by intertonguing into the Willow Canyon, Apache Canyon, and the lower 1,000 feet of the Shellenberger Canyon Formations; all of them are part of a conformable sequence above the limestone conglomerate to the south. Therefore, the upper contact of the Glance in the southern part of the range is placed at the top

of the highest conglomerate bed that is overlain by siltstone or sandstone that contains only a few thin beds of conglomerate, but in the area of intertonguing to the north, tongues of conglomerate are shown as Glance where they exceed 20 feet in thickness; otherwise, such tongues are arbitrarily not mapped separately.

#### WILLOW CANYON FORMATION

The Willow Canyon Formation was named by Tyrrell (1957) for exposures along Willow Canyon on the west flank of the Whetstone Mountains, and he described the type section in the SE¼ sec. 25, T. 18 S., R. 18 E. The upper member of Tyrrell's Willow Canyon Formation in the type section is similar to the upper 1,000 feet of the Willow Canyon Formation (restricted) in the Empire Mountains, where the formation is at least 3,000 feet thick. Because the lithology of the Willow Canyon Formation is more varied in the Empire Mountains, reference section there is designated that extends from the S½ sec. 18 to the NW¼ sec. 29, T. 18 S. R., 17 E. The reference section contains the lower two-thirds of the formation on the northwest side of a zone of normal strike faults and the upper three-fourths on the southeast side, and it shows the typical lithologic types of the formation.

The Willow Canyon Formation in the reference section consists predominantly of an alternating sequence of sandstone and siltstone that erodes to valleys and ridges of moderate relief. The sandstone is light yellowish gray, light pinkish gray, and light yellowish brown, arkosic, crossbedded, and locally conglomeratic. The pebbles and cobbles in the conglomeratic sandstone are well rounded and composed mainly of quartzite and some granodiorite and quartz diorite. The siltstone is commonly dark reddish brown in the lower two-thirds of the formation and olive gray to greenish gray in the upper one-third. A few thin beds of silty limestone and calcareous sandstone occur in the upper 300 feet of the formation, and a shaly siltstone about 250 feet below the top contains fossil gastropods of the genus Viviparus (W. A. Cobban, written commun., 1966), a long-ranging fresh-water form.

The contact with the overlying Apache Canyon Formation is placed arbitrarily at the horizon above which limestone is a dominant part of the lithology. Thus, some thin limestone beds are included in the Willow Canyon Formation. The Willow Canyon is at least 3,000 feet thick in the southern part of the Empire Mountains, but northward it grades laterally into the Glance Conglomerate by intertonguing and is absent along the north side of the range.

#### APACHE CANYON FORMATION

The Apache Canyon Formation was named by Tyrrell (1957) for exposures in Apache Canyon on the west flank of the Whetstone Mountains. He described the type section at that locality extending from the base in the SW¼NW¼ sec. 31, T. 18 S., R. 19 E., to the top in the NW¼NW¼ sec. 1, T. 19 S., R. 18 E.

The Apache Canvon Formation in the Empire Mountains and the western part of the Whetstone Mountains consists of thinly laminated to thick-bedded silty limestone, black shale, siltstone, and arkosic sandstone in an alternating sequence that erodes to low-rounded ridges and valleys. The limestone is generally dark gray to black and exhibits laminae as little as a millimeter thick of black limestone alternating with gray silty limestone that forms a distinctive lithologic type unlike any Paleozoic limestone in the region. The shales are dark-gray to black fissile shales intercalated with the limestone beds. Siltstone is dark gray to orange and red, calcareous, and thin bedded to massive. Massive red siltstone about 200 feet above the base contains a bed of gypsum about 2 feet thick, and some limestone beds above the gypsum contain numerous siltstone casts of what may have been gypsum crystals. The upper half of the Apache Canyon contains several beds of sandstone as much as 15 feet thick. The sandstone is yellowish gray, yellowish brown, and olive gray, fine to very coarse grained, partly crossbedded and partly massive, and arkosic. Biscuit-shaped bodies of silty limestone as much as 3 feet across and half a foot thick occur here and there in the limestone and shale beds.

The Apache Canyon Formation grades upward through increasing amounts of sandstone and siltstone into the Shellenberger Canyon Formation, the upper contact in the Empire Mountains being placed at the top of the uppermost limestone that is more than 3 feet thick and that is overlain by siltstone and sandstone containing only a few beds of limestone. The Apache Canyon Formation is at least 1,600 feet thick in the Empire Mountains but only 550 to 740 feet thick in the type area (Tyrrell, 1957, p. 101). Like the Willow Canyon Formation it intertongues northward with the Glance.

#### SHELLENBERGER CANYON FORMATION

The Shellenberger Canyon Formation was named by Tyrrell (1957) for exposures along the northeast-trending part of Shellenberger Canyon on the southwest flank of the Whetstone Mountains. Tyrrell (1957) described the type section of the Shellenberger Canyon Formation as extending from the base in the NE½NE½ sec. 34, T. 18 S., R. 18 E., to the top in the SW½NE½ sec. 4, T. 19 S., R. 18 E.

The Shellenberger Canvon Formation in the Empire and Whetstone Mountains consists of an alternating sequence of shale, siltstone, sandstone, and a little limestone that erodes to valleys and ridges of moderate relief. The shale and siltstone are commonly shades of olive brown, olive gray, and greenish gray although in a few places they are reddish brown. Small bits of carbonized plant fragments are present in many of the somber-hued shales and siltstones, and silicified logs as much as 5 feet in diameter abound in some beds. The sandstone is olive brown, olive green, olive gray, and pinkish gray, fine to very coarse grained, arkosic, massive to crossbedded and lenticular. Some of the sandstone beds are conglomeratic and contain pebbles of quartzite and chert as much as an inch in diameter. One sandstone bed about 500 to 800 feet below the top of the formation has a distinctive basal conglomerate as much as a foot thick that contains well-rounded pebbles of black and red chert. The siltstone above this sandstone contains a calcareous zone about 2 to 10 feet thick that contains scattered well-rounded pebbles and cobbles of limestone and quartzite.

Limestone is mainly restricted to the lower 1,300 feet of the formation; some of it is laminated like that in the Apache Canyon Formation and some is thin bedded. Two beds about 1,000 and 1,300 feet above the base are composed of pelecypod and gastropod shells in a matrix of calcite, sand, and silt. Fossils of an Early Cretaceous dinosaur (Miller, 1964, p. 378) were obtained by Moore and Miller (1960, p. 59-60) from a shaly bed between the shell beds in the Empire Mountains, and a fossil clam *Trigonia* n. sp., was found in calcareous siltstone that overlies a similar shell bed in the Whetstone Mountains (Tyrrell, 1957).

The contact of the Shellenberger Canyon Formation with the overlying Turney Ranch Formation in both ranges is placed at the top of a light-brown sandstone that is overlain by red shale and siltstone that contains nodules of gray limestone and that is overlain in turn by light yellowish-gray to light pinkish-gray sandstone. The Shellenberger Canyon generally ranges in thickness from 3,950 to 4,330 feet, but to the north, the lower 1,000 feet of the formation grades into the Glance Conglomerate.

#### TURNEY RANCH FORMATION

The Turney Ranch Formation was named by Tyrrell (1957) for its typical exposures immediately north and west of Turney Ranch in the Whetstone Mountains. This ranch is now known as the Clyne Ranch and is located in secs. 21, 22, and 27, T. 19 S., R. 18 E., just inside the southeast corner of the Empire Mountains quadrangle. The

described type section extends from the base in the SE¼NW¼ sec. 11, T. 19 S., R. 18 E., to the top in the SW¼ sec. 15, T. 19 S., R. 18 E.

The Turney Ranch Formation in the Empire and Whetstone Mountains consists of red siltstone and shale alternating with light-pinkish-gray sandstone. It erodes to valleys and ridges of moderate relief. The siltstone and shale are various shades of red and purplish red though locally gray near nodules of limestone. Some beds are calcareous. In places, gradations from red fine-grained sandstone into siltstone and finally into shale occur both vertically and laterally. The sandstone is light pinkish gray to pale orange, medium to coarse grained, arkosic, and crossbedded. Lenses of pebble conglomerate containing chert and quartzite pebbles set in an arkosic matrix occur here and there in the sandstone beds. Scour-and-fill structures a few inches deep occur at the base of some of the sandstone beds.

In the type area, the Turney Ranch Formation is overlain with angular unconformity by upper Cenozoic gravel deposits. In the Empire Mountains, tilted and faulted Turney Ranch beds are locally overlain with angular discordance by conglomerate and andesitic breccia that are probably of Late Cretaceous age, for they are very similar to the lower part of the Salero Formation which has been assigned a Late Cretaceous age by Drewes (1968) on the basis of potassium-argon age determinations of the upper part of the formation. The Turney Ranch Formation is at least 3,200 feet thick in the type area. Its original thickness cannot be determined because its top is eroded.

#### AGE

The formations of the Bisbee Group in the Empire and Whetstone Mountains are all part of a conformable sequence and are considered to be of the same general age. The only fossils that are diagnostic as to age are of an Early Cretaceous dinosaur and of the mollusk Trigonia, all from the Shellenberger Canyon Formation near the middle of the sequence. Tyrrell (1957) submitted specimens of Trigonia to Prof. A. A. Stovanow, who determined that they are a new species that is not represented in the fauna from the Bisbee Group in the type area. However, Stovanow considered that the beds containing Trigonia in the Whetstone Mountains are probably no older than the upper Trinity Group (Upper Cretaceous) and that they were deposited in the same sea as the Bisbee Group in the Bisbee area. The Bisbee Group in the Empire Mountains quadrangle rests unconformably on a variety of older rocks, the youngest of them being sedimentary and volcanic rocks that are probably equivalent to the Gardner Canyon Formation of Triassic age (Drewes, 1968). The Bisbee is overlain with a strong angular unconformity by sedimentary and volcanic rocks that are similar to the Salero Formation of Late Cretaceous

age. These geologic relations, together with the meager fossil evidence, indicate that the Bisbee Group in this area was probably deposited during Early Cretaceous time.

### PANTANO FORMATION

BY TOMMY L. FINNELL

The Pantano Formation, herein adopted, was introduced by C. F. Tolman in unpublished data on the geology of the Tucson 30-minute quadrangle prepared for the U.S. Geological Survey about 1912. The first published detailed description of the Pantano, other than a brief reference by King (1939, p. 1692), was by Brennan (1962, p. 46-53). In this report Brennan selected and described the type section along the south side of U.S. Highway 80 (now Interstate 10) extending from about the east line of sec. 2, T. 17 S., R. 17 E. (top of section), westward along the highway to the center of the W½ sec. 31, T. 16 S., R. 17 E., in the Empire Mountains quadrangle, southeastern Arizona. Detailed mapping has revealed that several normal faults repeat parts of the Pantano Formation in Brennan's type section. Also at least three angular unconformities within the formation omit certain stratigraphic units from the type section.

The Pantano Formation is readily divisible into five units that have a combined thickness of at least 6,400 feet:

Unit 1, the lowest unit, is faulted against vertical beds of the Bisbee Formation of Early Cretaceous age along Interstate 10, but it lies on eroded edges of Tertiary or Cretaceous volcanic and sedimentary rocks in Barrel Canyon near State Highway 83 about 11 miles south-southwest of the town of Pantano. It consists of reddish-brown bouldery mudstone and siltstone, conglomeratic sandstone, conglomerate, and at least one rhyolite tuff of probable ash-flow origin. The sandstone and conglomerate are mostly in poorly defined layers, but in places they form alternating beds as much as 2 feet thick. This unit is at least 1,600 feet thick.

Unit 2 is red and yellowish-gray mudstone alternating with fine conglomerate and conglomeratic sandstone of the same colors; in the upper 100 feet it contains at least four beds of olive-gray partly onlitic sandy argillaceous limestone. The whole unit is about 500 feet thick.

Unit 3 is massive to poorly bedded reddish-brown conglomerate with a few mudstone and coarse-grained sandstone partings. It pinches out against the lower units a few hundred feet south of the type section but extends northward for at least 2 miles. It is at least 2,500 feet thick.

Unit 4 is a flow of dark-purplish-gray andesite called Turkey-Track Porphyry by Cooper (1961) because of the large plagioclase phenocrysts. The upper part of the flow is eroded, and it thins to a featheredge about 300 feet south of the highway.

Unit 5 locally contains a dark-purplish-gray volcanic conglomerate at the base. Where present, this grades up into purplish-gray and yellowish-brown conglomerate and coarse conglomeratic sandstone that grade upward and northward into fine-grained tuffaceous sandstone, siltstone and gypsiferous claystone containing a few interbeds of sandstone, and angular conglomerate or monolithologic breccia.

At least three tuff beds occur near the middle. The unit is at least 1,800 feet thick.

The subangular to well-rounded pebbles, cobbles, and boulders in the Pantano Formation were mainly derived from the Bisbee Formation, but partly from the Precambrian, Paleozoic, lower Tertiary(?), and Upper Cretaceous rocks of the region. In unit 3, boulders and cobbles of rhyodacite welded tuff identical in appearance to an Upper Cretaceous welded tuff are a conspicuous but minor constituent.

The surface upon which the Pantano Formation was deposited had considerable relief, as indicated by apparent depositional thinning and pinch-out of the lower units. Local deformation during deposition is suggested by angular unconformities between some of the units. For example, unit 5 rests on the andesite flow (unit 4) along the highway, but about half a mile to the south, it rests unconformably on limestone of unit 2.

In places, faulted and tilted beds of the Pantano are overlain by generally flat-lying gravel deposits of Pliocene(?) and Pleistocene age. The Pliocene(?) gravels are light pinkish gray to light yellowish gray, and they may have covered the entire type area of the Pantano Formation at some earlier time for remnants of these gravels are known to overlie the upper four units.

No indigenous fossils have been found in the Pantano Formation, although its Paleozoic and Lower Cretaceous clasts are fossiliferous. However, its age can be deduced from radiometric dating. Dates on minerals from the rhyolite tuff of unit 1 have been determined by Damon and Bikerman (1964, p. 69) as follows:

Sanidine (K-Ar)  $36.7 \pm 1.1$  m.y.

Biotite (K-Ar)  $32.8 \pm 2.7$  m.y.

P. E. Damon (written commun. 1966) determined the radiometric age of plagioclase from the andesite flow to be 24.4±2.6 m.y. (K-Ar). On this basis, the Pantano Formation seems to range in age from early Oligocene to early Miocene.

# BARDSTOWN MEMBER OF THE DRAKES FORMATION IN CENTRAL KENTUCKY

By WARREN L. PETERSON

Prepared in cooperation with the Kentucky Geological Survey

The Drakes Formation of Late Ordovician age consists of two members at its type locality in south-central Kentucky: the Rowland Member and the overlying Preachersville Member (Weir and others, 1965). In the vicinity of Bardstown, Ky., the Drakes Formation is divided into three members: the lowest, the Rowland Member, is closely similar to the Rowland in south-central Kentucky; the highest is the Saluda Dolomite Member; the middle member is herein named the Bardstown Member for exposures in the vicinity of Bardstown, Nelson County, Ky. The type section (fig. 3) is in Nelson County, Ky., in the southwest quarter of the Maud quadrangle in a roadcut along U.S. Highway 150, 1.35 miles northwest of Fredericktown, Washington County, Ky., and 0.95 mile from the west border of the quadrangle (lat 37°46′20″ N.; long 85°21′30″ W.).

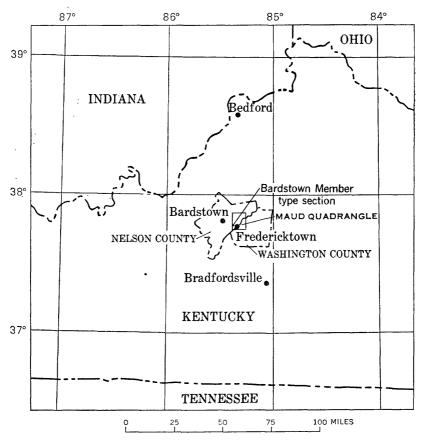


FIGURE 3.—Location of the type section of the Bardstown Member of the Drakes Formation.

About 90 percent of the Bardstown Member consists of fine- to medium-grained, gray to greenish-gray muddy limestone that contains scattered to abundant fossils and fossil fragments. Weathered rock is yellowish gray with a shaly aspect. Samples of the matrix yielded 27 to 50 percent insoluble residue. Beds are discontinuous and lentic-

ular and commonly 1 to 8 inches thick. About 10 percent of the member is gray bioclastic and coquinoidal limestone composed of whole fossils and fossil fragments in a fine-grained to very coarse grained matrix cemented in part by sparry calcite. The matrix is predominantly calcitic fossil debris, with variable amounts of non-carbonate silt and clay. It occurs commonly in ½- to 3-inch-thick discontinuous beds, lenses, and knots with rough to smooth surfaces that form thin ledges in gully exposures or slump as resistant blocks. A small percentage of the member is calcareous shale (containing more than 50 percent insoluble residue) that is physically indistinguishable from the muddy limestone.

Fossils in the Bardstown Member are principally brachiopods, bryozoans, pelecypods, gastropods, horn corals, and colonial corals. Horn corals ("Streptelasma") are sparse to abundant and, in the vicinity of Bardstown, are probably restricted to this member. The colonial corals are Tetradium (with tiny rectangular corallites) and Foerstephyllum, Favistella, and Caleopecia (with larger hexagonal corallites) (Browne, 1964). The corals with hexagonal corallites occur in heads as much as 4 feet across. The colonial coral heads are concentrated in two to four layers in the middle two-thirds of the unit. Locally, layers as much as 4 feet thick are 25 to 50 percent coral heads. The coral layers in the Bardstown Member are probably correlatives of the Otter Creek coral bed, which occurs at the base of the Preachersville Member of the Drakes Formation in eastcentral Kentucky (Simmons and Oliver, 1967). Massive coral layers in some exposures were called the "Bardstown coral reef" by Foerste (1903). Fossils in the Bardstown Member have been listed by Butts (1915), Browne (1964, 1965), and Hatfield (1968) under different formation names discussed in another part of this paper.

The top of the Bardstown Member is placed where the muddy fossiliferous limestone grades abruptly into dolomite of the overlying Saluda Dolomite Member. The contact is not obvious but can be located by application of dilute hydrochloric acid. The contact at the base of the Bardstown Member is commonly sharp or gradational through a few inches. In some places, however, the basal contact is gradational by interlayering through an interval up to 4 feet thick. The contact is placed at the base of rock with abundant fossils.

The Bardstown Member is commonly 25 to 35 feet thick but ranges in thickness from 12 to 40 feet within a 15-mile radius of the type section. It is recognizable along the outcrop northward to near Bedford, Trimble County, Ky., and southward and southeastward at least as far as Bradfordsville, Marion County, Ky.

The Bardstown Member has previously been called the unnamed member of the Drakes Formation by Peterson (1966, 1967, 1968).

Approximately the same unit was designated the Liberty Limestone by Nosow (1959, fig. 10), the Liberty Formation by Butts (1915), the Liberty Formation and lower part of the Whitewater Formation by Browne (1964), and part of the Tanners Creek Formation by Hatfield (1968). The Liberty Limestone (or Formation) which was named in Indiana (Nickels, 1903) has been recognized in Kentucky only on a faunal basis, and the name is, therefore, considered inappropriate for a rock-stratigraphic unit in Kentucky. The name Tanners Creek Formation defined by Fox (1962) at Tanners Creek, Dearborn County, Ind., was rejected for use in Indiana by Brown and Lineback (1966) because of the lack of a regionally traceable base. The name is not considered usable in Kentucky for the same reason. The rocks of the Bardstown Member have been considered to be of Richmond (Late Ordovician) age by all writers.

Drakes Formation at type section of the Bardstown Member of the Drakes Formation

[Section measured in the southwest quarter of the Maud quadrangle, Nelson County, in roadcut along U.S. Highway 150, 1.35 miles northeast of Fredericktown, Washington County, Ky., lat 37°46′20″ N.; long 85°21′30″ W.]

Brassfield Dolomite (basal part only):	Thickness
22. Dolomite, light-gray to light-brown, fine- to medium-grained	(feet) 4+
Drakes Formation:	•
Saluda Dolomite Member:	
21. Shale, grayish-green, silty and clayey; contains no mega-	
	2, 3-2, 8
20. Dolomite similar to unit 19 with abundant Tetradium	
heads; in part recrystallized, flattened parallel to	
beddingbedding	4, 2-4, 7
19. Dolomite, greenish-gray; color banded parallel to bedding	
on surface into 1 to 2 inch bands of yellowish brown,	
reddish brown, dark gray, and grayish green; banding	
probably reflects faint bedding; very fine grained; slightly	
silty and clayey; unfossiliferous except for poorly pre-	
served dolomitized cylindrical bryozoans; some recrystal-	
lized calcitic bryozoans in upper 2 feet; dolomite pebbles	
in upper 2 feet; worm bored throughout; forms massive,	
jutting rounded ledge	6. 3
18. Dolomite, greenish-gray, very fine grained, massive; less	
resistant than overlying dolomite, silty and clayey;	
contains no megafossils; worm bored in lower 3 feet;	
upper 3 to 6 inches slightly indented and less resistant.	
upper 5 to 6 menes singulary indented and less resistant.	1. 5
Total Saluda Dolomite Member	20. 6
Bardstown Member:	
17. Limestone, greenish-gray, muddy, fine-grained; contains	

17. Limestone, greenish-gray, muddy, fine-grained; contains scattered to abundant fossils and fossil fragments; in faint beds about 2 inches thick; worm bored; contains sparse thin beds, lenses, and knots of purer, more re-

Drakes Formation—Continued	Thickness
Bardstown Member—Continued sistant gray fossiliferous limestone; fossils include brachiopods, cylindrical bryozoans, gastropods, pelecy-	(feet)
pods, and horn coral	7. 6
occur closely packed in parts of unit and sparsely in other parts	4. 3
15. Limestone, greenish-gray, muddy, fine-grained; contains sparse to abundant fossils and fossil fragments; in beds 1 to 5 inches thick; less abundant but more resistant limestone makes up about 20 percent of rock, is medium gray, has medium- to coarse-grained matrix and abundant fossils and fossil fragments in beds, lenses, and knots as much as 2 inches thick; minor olive-gray shale occurs interbedded with limestone; all gradations occur from nearly pure limestone to calcareous shale. Fossils include brachiopods, cylindrical bryozoans, horn corals, gastropods, pelecypods, and scattered colonial coral	
heads	7. 2
sparse to abundant	
13. Limestone and minor shale; similar to unit 15	T. 1-0. 0
to both Bardstown and Rowland lithologies	3. 0
Total Bardstown Member	29. 3
Rowland Member:  11. Limestone, greenish-gray, medium grained; abundant small fossil fragments; single bed faintly parted into 1-inch layers	1. 5
10. Limestone, greenish-gray to grayish-green, very fine grained, silty and clayey, dolomitic, thinly laminated; parted into 1- to 2-inch layers; contains no megafossils; weakly resistant; weathers with shaly aspect; some mud-	10.0
9. Limestone, greenish-gray; similar to unit 10 but contains scattered to abundant fossil fragments and ostracode shells; in beds 1 to 11 inches thick; no conspicuous lami-	16. 8
nations; forms resistant outcropping ledge  8. Limestone; similar to unit 10	5. 9 12. 2
7. Shale, olive-gray, calcareous, silty, clayey, weakly fissile; contains no megafossils	. 3

Drakes Formation—Continued	Oth tales as a
Rowland Member—Continued	Thickness (feet)
6. Limestone, greenish-gray, impure, very fine grained, lami-	• ,
nated; parts into ½- to 1½-inch layers; contains some	
fossil fragments in upper part; more resistant to weath-	
ering than units such as 10	4. 9
5. Limestone similar to unit 4 in beds as much as 1 inch thick	
and limestone similar to unit 3 mostly as thin partings	1. 0
4. Limestone, light-olive-gray, fine- to medium-grained;	
contains pebbles of greenish-gray limestone; single bed	
with smooth top and bottom; contains no megafossils	. 3
3. Limestone, grayish-green; conspicuously green on outcrop,	
silty and clayey, very fine grained; contains sparse small	
brachiopods, bryozoans, and fossil fragments; weathers	
shaly	1. 8
2. Shale, dark-olive-gray, calcareous, clayey, silty, weakly	
fissile; contains no megafossils.	. 3
Total Rowland Member	45. 0
Grant Lake Limestone:	
1. Limestone and shale (not measured).	

# THE CANTWELL FORMATION OF THE CENTRAL ALASKA RANGE

By Jack A. Wolfe and Clyde Wahrhaftig

#### NAME AND TYPE AREA

The Cantwell Formation was first described by Eldridge (1900, p. 16) as the Cantwell Conglomerate from exposures in the canyon of the Nenana River, then called the Cantwell River. Its type locality is considered to be the east wall of the canyon of the Nenana River from the mouth of Slime Creek northward for about 7 or 8 miles in the Healy C-4 quadrangle (fig. 4, loc. 1). This type locality agrees with Eldridge's original text (1900, p. 1) but not with his map (map 3) nor with the statement in Wahrhaftig (1958, p. 8), both of which are in error.

## DISTRIBUTION

The Cantwell Formation occupies a large synclinorium extending along the center of the Alaska Range from the headwaters of the Wood River westward to the Muldrow Glacier (fig. 4). Brooks and Prindle (in Brooks, 1911) and Reed (1961) have mapped bodies of the Cantwell farther southwest along the north base of the range. Capps (1933) mapped a large area of the Cantwell Formation on the south side of the Denali fault between Foggy Pass and Anderson Pass, an area separated from the main body of the Cantwell by a belt of Paleozoic and Triassic rocks along the north side of the fault.

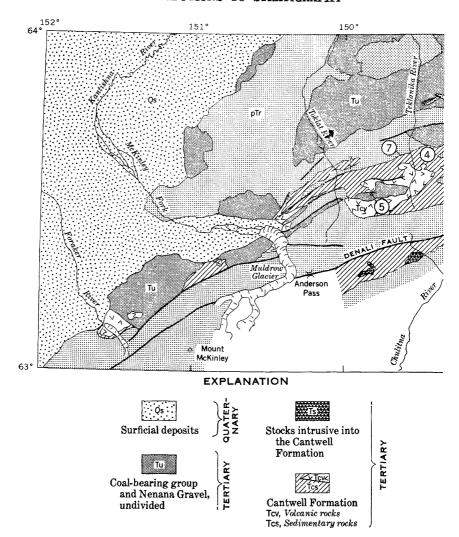
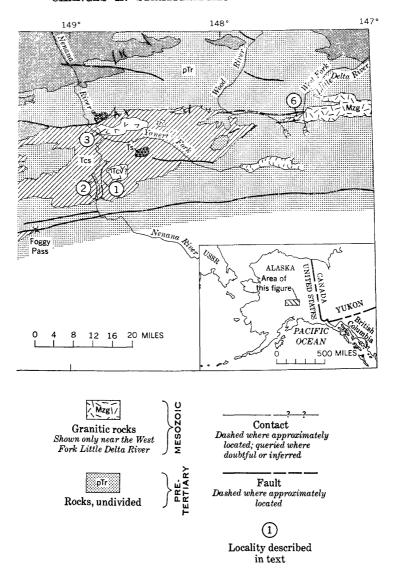


FIGURE 4.—Geologic sketch map of the central Alaska Range showing distributext. Geology modified from Reed (1961, pl. 1),



tion of the Cantwell Formation, related formations, and localities referred to in Wahrhaftig (1958, pl. 1), and Capps (1912, pl. 2).

#### LITHOLOGY

The Cantwell Formation consists predominantly of interbedded conglomerate, sandstone, argillite, shale, and coal but also contains volcanic rocks, especially near its top. The Cantwell is intruded by an abundance of sills and dikes ranging in composition from diabase to rhyolite and by monzonite stocks as much as 3 to 4 miles across.

At two localities the Cantwell Formation is clearly younger than large intrusive bodies. (1) At the pass between the headwaters of the Wood River and the West Fork Little Delta River (fig. 4, loc. 6), it rests unconformably on granodiorite and quartz monzonite at the west end of a large batholith which was traced eastward by Capps (1912, pl. 2) as far as Delta Creek, a distance of 30 miles; the granodiorite beneath the contact has been weathered to grus for a thickness of 20 feet. (2) On the east side of the Muldrow Glacier, dikes believed to be feeders to the volcanic rocks of the Cantwell Formation cut the intrusive body underlying Mount Eielson (Reed, 1933).

The Cantwell Formation is generally moderately well consolidated and is locally very well indurated. Some beds in the formation are extremely well sorted pebble conglomerate consisting of quartz, chert, quartzite, and argillite pebbles with little or no matrix. Few of the pebbles are less than half, or more than twice, the median size. The pebbles are indented and moulded against each other, possibly as a result of solution and redeposition of silica under tectonic pressure; the conglomerate therefore has negligible porosity.

Coal in the Cantwell Formation is generally of bituminous rank, even where beds are tightly folded and the conglomerate compressed, as in the vicinity of mafic dikes.

Dark flows and rhyolite tuffs occupy an open syncline centered on Mount Fellows (fig. 4, loc. 3). A belt of predominantly silicic volcanic rocks is exposed at the top of the formation from Double Mountain (fig. 4, loc. 4) westward to the Muldrow Glacier. These silicic volcanic rocks give Polychrome Pass (fig. 4, loc. 5) its color variety. In both the Mount Fellows and Polychrome Pass areas, the volcanic rocks are at the top of the section, yet only 1,000 to 3,000 feet of sedimentary rocks lies between the volcanic rocks and the base of the formation. The volcanic rocks at Mount Fellows and Polychrome Pass may be equivalent to lower parts of the Cantwell Formation at other places, or they may have erupted in areas where highlands in the pre-Cantwell topography rose a few thousand feet above the base of the formation elsewhere.

The lithology of the clasts in the Cantwell Formation varies from place to place, suggesting that the formation was derived from at least three different sources. In the type area along the Nenana River, the sandstone and conglomerate are predominantly dark gray and con-

sist largely of argillite, chert, quartzite, and quartz grains and pebbles. The source of this dark-gray facies was probably south of the Alaska Range, possibly in the Mesozoic rocks in the Talkeetna Mountains or the southern Alaska Range. Along the north edge of the area of outcrop of the Cantwell east of the Nenana River, this facies interfingers with light-brown to white sandstone and conglomerate consisting largely of schist fragments derived from the Birch Creek and Totatlanika Schists immediately to the north. From the vicinity of Polychrome Pass (fig. 4, loc. 5) westward, the dark-colored Cantwell is replaced by light-brown conglomerate and sandstone that consist largely of clasts of gray limestone in a light-brown matrix of unknown composition and origin. The westward limit of this light-colored limestone-bearing facies is unknown. Its source could have been the area of Devonian limestone near the crest of the range (and on the north side of the Denali fault) immediately south of its area of outcrop.

The limestone-bearing facies has not been mapped in the Cantwell Formation south of the Denali fault, as this body has not been visited since Capps' reconnaissance in 1930. A study of this facies and its correlation with facies north of the fault might give information on the lateral displacement of the fault since Paleocene time.

### THICKNESS

The total thickness of the Cantwell Formation, whose upper surface has been eroded, is unknown. Generally, the thickness ranges from 2,000 to 5,000 feet. The maximum thickness is about 10,000 feet in a reference section (fig. 4, loc. 2) in the walls of the Nenana Canyon between Clear Creek and Carlo (Wahrhaftig, 1958, pl. 3) but may include an unknown thickness of sills.

#### AGE

The Cantwell Formation has had a varied history of age assignments. Eldridge (1900, p. 16), when first describing the Cantwell, was unable to assign an age to the formation. Brooks (1911, p. 78–83), who, with Prindle, mapped the Cantwell Formation from Mount McKinley to the Nenana River (fig. 4), correlated it on the basis of lithologic similarity with the Nation River Formation, then considered to be of Carboniferous age; but Brooks, in making this correlation had to explain away as faulted inliers beds containing plant remains which he had collected and which F. H. Knowlton reported to be Tertiary.

Moffit (1915) found plant fossils in rocks about 15 miles east of the Nenana that are similar to those of the type area. Knowlton and Hollick (in Moffit, 1915, p. 48) assigned these fossils to the Eocene but noted that Brooks' collections were similar. For 20 years there-

after, geologists working in Alaska were troubled by the fact that both the Cantwell Formation and an apparently much younger coal-bearing formation of the Nenana coal field to the north contained plants of supposed Eocene age.

In 1936, Ralph Chaney made new collections in the Cantwell at another highly fossiliferous locality and determined the fossils to be Cretaceous in age (Chaney, 1937), the age for the Cantwell quoted by Capps (1940, p. 118). In a further refinement of its age, Imlay and Reeside (1954, p. 235), noting the similarity of the plant species identified from the Cantwell by Chaney to those in the Upper Cretaceous Chignik Formation of the Alaska Peninsula and the Lower Cretaceous Melozi and Kaltag Formations of western Alaska, placed the Cantwell in the Lower Cretaceous Albian Stage.

Fossil plants have been collected from several localities in the Cantwell Formation, but only one locality, that of Chaney's collections, has furnished well-preserved and abundant material. Reexamination of Chaney's collections (Univ. of California Museum Paleont. loc. P3654—loc. 7 of fig. 4) indicates that the following taxa are present:

Glyptostrobus? sp.

Metasequoia occidentalis (Newb.) Chan.

Sparganium antiquum (Newb.) Berry

Planera microphylla Newb.

Cocculus flabella (Newb.) Wolfe

Cissus sp. aff. C. marginata (Lesq.) R. W. Br.

Grewiopsis auriculaecordatus (Holl.) Wolfe

Dicotylophyllum flexuosa (Newb.) Wolfe

All these named species are indicators of a Tertiary age; Cissus marginata is known from both Upper Cretaceous and Paleocene rocks (Brown, 1962). The species of Planera, Cocculus, Grewiopsis, and Dicotylophyllum were recorded from the Chickaloon Formation of the Cook Inlet region (Wolfe and others, 1966) and from other rocks thought or known to be of Paleocene age. Inasmuch as the particular species of these four genera have not been found in rocks younger than Paleocene, we therefore consider the known fossil plants to be indicative of a Paleocene age for the Cantwell Formation.

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